

UNITED STATES DEPARTMENT OF THE INTERIOR
GEOLOGICAL SURVEY

Preliminary assessment of the geochemistry and mineral favorability of the
postorogenic granites of the southeastern Arabian Shield, Kingdom of Saudi Arabia
by
J. S. Stuckless 1/, G. VanTrump, Jr. 1/, E. H. Christiansen 2/, C. A. Bush 1/,
C. M. Bunker 1/, and A. J. Bartel 1/

Open-File Report 83- 486

Prepared for the Ministry of Petroleum and Mineral Resources
Deputy Ministry for Mineral Resources
Jiddah, Kingdom of Saudi Arabia

This report is preliminary and has not been reviewed for conformity with
U.S. Geological Survey editorial standards and stratigraphic nomenclature.

1/ U.S. Geological Survey, Denver, CO 80225
2/ Department of Geology, University of Iowa,
Iowa City, Iowa 55242

CONTENTS

| | <u>Page</u> |
|-----------------------------------|-------------|
| ABSTRACT..... | 1 |
| INTRODUCTION..... | 2 |
| ANALYTICAL PROCEDURES..... | 4 |
| RESULTS AND DISCUSSION..... | 18 |
| Petrogenetic considerations..... | 18 |
| Economic considerations..... | 22 |
| Suggestions for further work..... | 36 |
| SUMMARY AND CONCLUSIONS..... | 37 |
| DATA STORAGE..... | 37 |
| REFERENCES CITED..... | 38 |

ILLUSTRATIONS

| | |
|------------------------------------------------------------------------------------------------------------------------------------------|----|
| Figure 1. Location map for plutons sampled in the southeastern part of the Arabian Shield..... | 3 |
| 2. Ternary diagram showing normative feldspar data for granite samples from the southeastern Arabian Shield..... | 23 |
| 3. Ternary diagram showing normative quartz, albite, and orthoclase data for granite samples from the southeastern Arabian Shield..... | 24 |
| 4. Map showing areal distribution of uranium values..... | 28 |
| 5. Map showing areal distribution of thorium values..... | 30 |
| 6. Diagram showing plot of uranium versus radium equivalent uranium values for granite samples from the southeastern Arabian Shield..... | 33 |
| 7. Map showing areal distribution of Th/U ratios..... | 34 |

TABLES

Page

| | |
|---------------------------------------------------------------------------------------------------------------------------------------------------|----|
| Table 1. Chemical and normative compositions of granitic samples from the southeastern Arabian Shield..... | 5 |
| 2. Trace-element concentrations in granite samples from the southeastern Arabian Shield..... | 12 |
| 3. Statistical summary for chemical compositions of granite samples from the southeastern Arabian Shield..... | 19 |
| 4. Comparison of means and standard deviations for subgroupings of granite samples from the southeastern part of the Arabian Shield..... | 21 |
| 5. Correlation matrix for postorogenic granite samples from the southeastern Arabian Shield..... | 25 |

PRELIMINARY ASSESSMENT OF THE GEOCHEMISTRY AND
MINERAL FAVORABILITY OF THE POSTOROGENIC GRANITES
OF THE SOUTHEASTERN ARABIAN SHIELD,
KINGDOM OF SAUDI ARABIA

by

J. S. Stuckless^{1/}, G. VanTrump, Jr.^{1/},
E. H. Christiansen^{2/}, C. A. Bush^{1/},
C. M. Bunker^{1/}, and A. J. Bartel^{1/}

ABSTRACT

Chemical analyses of samples for 19 postorogenic plutons from the southeastern Arabian Shield show that these rocks have average potassium/rubidium ratios (162) and average rubidium/strontium ratios (11.8) characteristic of highly evolved granites. Most of the analyzed samples are peraluminous. Three plutons are physically similar in terms of shape and megascopic textural zonation to peralkaline complexes in the northeastern part of the Shield, but none of the samples from these plutons is peralkaline. However, these plutons do contain the least-evolved samples.

Zinc, yttrium, uranium, thorium, and possibly copper, each occur in anomalously high concentrations in at least one pluton relative to contents typically cited for granite. The average regional concentrations of copper and zinc are anomalously high. These facts suggest at least a moderate potential for mineralization in the southeastern part of the Shield. Good correlations (r of 0.5 to 0.8) between uranium, thorium, yttrium, and rubidium and an excellent correlation ($r=0.98$) between uranium and radium-equivalent uranium suggest that secondary deposits of these elements are unlikely and that magmatic deposits, especially of a pegmatitic nature, are more likely.

Concentrations of some major elements and several trace elements covary with sample location as measured by latitude and longitude. This fact is interpreted to reflect regional variations in the protolith that have been proposed on the basis of lead isotopic data. The chemical variations suggest a largely oceanic crustal component for granites in the southwestern part of the southeastern Shield and a largely continental crustal component to the northeast.

^{1/} U.S. Geological Survey, Denver, Colorado 80225

^{2/} Department of Geology, University of Iowa,
Iowa City, Iowa 55242

Further quantitative analysis is recommended to verify and accurately delineate anomalous concentrations of trace elements. Trend surface analysis of all available data for postorogenic granites from the eastern Arabian Shield is suggested as a method for testing regional variations in the protolith that might control locations of highly favorable areas of ore deposition.

INTRODUCTION

This report is the result of research performed by the U.S. Geological Survey (USGS) in accordance with a work agreement with the Ministry of Petroleum and Mineral Resources, Kingdom of Saudi Arabia. The research is part of a study of the petrogenesis and mineral potential of the granitic rock of the Arabian Shield. This report provides reconnaissance results for postorogenic granites within a region (fig. 1) identified as anomalously metalliferous by du Bray and others (1983).

Postorogenic granites were intruded throughout the Arabian Shield from about 650 m.y. ago until at least 540 m.y. ago (Fleck and others, 1980). The granites are typically hypersolvus members of the calc-alkaline suite (Stoeser and Elliott, 1980), although subsolvus examples are common, particularly in the southeastern part of the Shield (du Bray and others, 1983). Chemically, the granites have high silica contents and range from peralkaline to peraluminous (Stoeser and Elliott, 1980; Stuckless and others, *in press*).

Postorogenic granites were chosen for this investigation because of their probable association with ore deposits in other parts of the world (for example, Wilson and Åkerblom, 1980; Doe and others, *in press*) and because of their importance to investigations of crustal evolution (for example, Anderson and others, 1980; Anderson, 1983). Many of these granites are highly evolved petrologically (for example, Barker and others, 1976; Anderson and Cullers, 1978; Cullers and others, 1981), and, in the area of the Arabian Shield, they contain anomalous amounts of incompatible trace elements (Harris and Marriner, 1980; Cole and others, 1981; Radain and others, 1981).

The classification of igneous rocks used in this report is that recommended by the IUGS Subcommission on the nomenclature of plutonic rocks (Streckeisen, 1973). The granites are subdivided on the basis of alumina saturation as defined by Shand (1951). Rocks with molar ratios of $Al/(Na+K) < 1$ are peralkaline, $Al/(Na+K) > 1$ and $Al/(Na+K+Ca) < 1$ are metaluminous, and $Al/(Na+K+Ca) > 1$ are peraluminous.

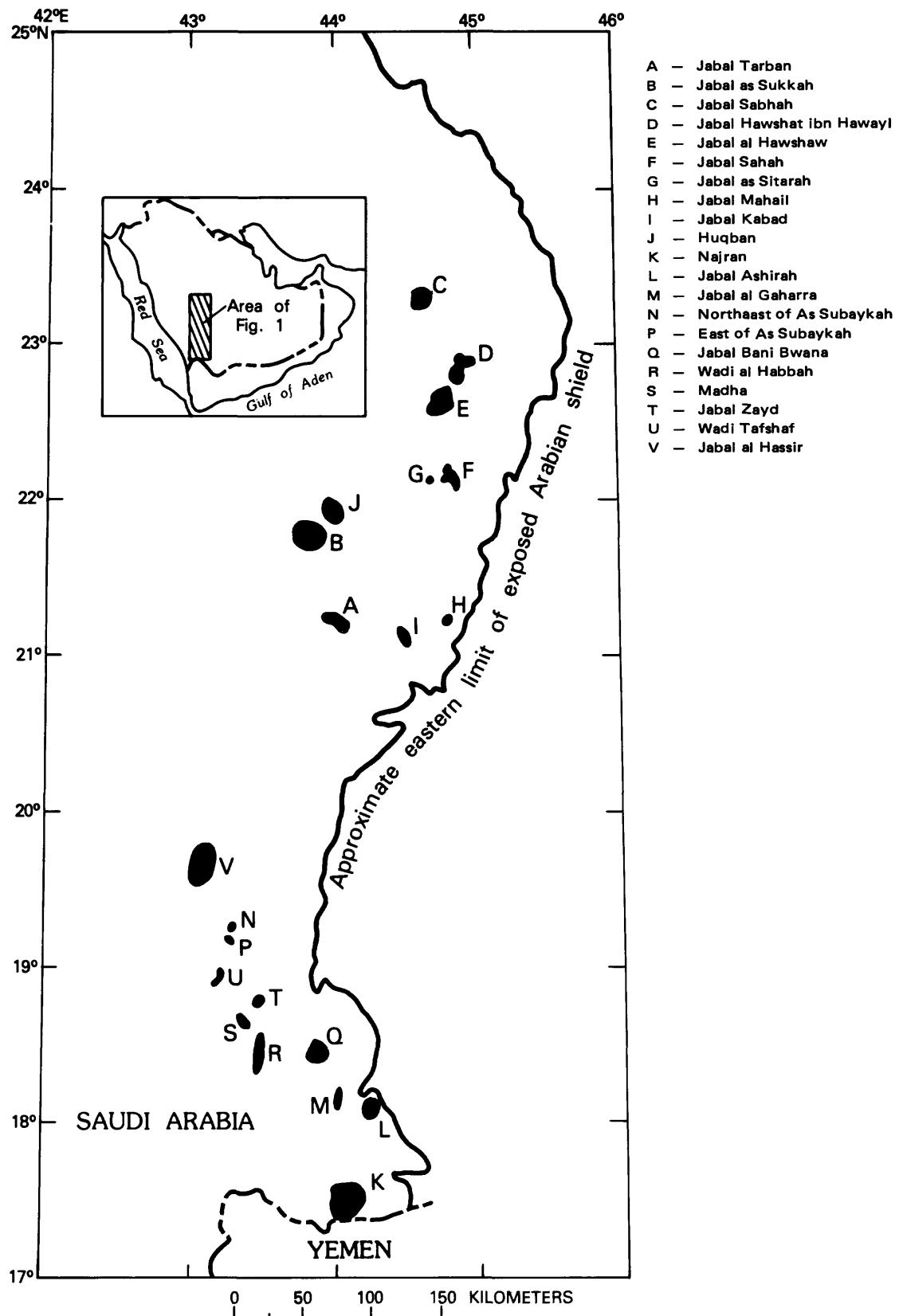


Figure 1.--Location map for plutons sampled in the southeastern part of the Arabian Shield. Letter designations correspond to symbol designations on table 1 and to location names on table 2.

ANALYTICAL PROCEDURES

Concentrations of major elements (table 1) were obtained by high-precision X-ray fluorescence (Taggart and others, 1982). This technique uses an 0.8 g aliquot of fused sample powder. Results are considered to be precise and accurate within \pm 2 percent of the amount reported (2σ) for abundances in excess of 1 percent absolute. A fourth digit is reported for SiO_2 and Al_2O_3 , which is not significant for any one sample but which may be significant in the statistical treatment of the entire data set (A. T. Miesch, oral commun., 1980). Values for MgO , TiO_2 , P_2O_5 , and MnO that are less than the detection limit (0.10, 0.02, 0.05, and 0.02, respectively) are arbitrarily assigned values of one-half the value of the detection limit for purposes of normative and statistical calculations.

Uranium and thorium concentrations (table 2) were determined on 8 to 10 g aliquots of sample powder by the delayed neutron technique (Millard, 1976). Accuracies for reported values, based on counting statistics, vary with concentration. Uranium concentrations are accurate to between \pm 4 and \pm 8 percent for more than 2 parts per million (ppm), to between \pm 8 and \pm 14 percent for 1 to 2 ppm, and to between \pm 14 to \pm 36 percent for less than 1 ppm. Thorium concentrations are accurate to between \pm 8 and \pm 30 percent for more than 6 ppm and to about \pm 50 percent for less than 6 ppm (2σ). Thorium values greater than 10 ppm are generally accurate to within \pm 15 percent of the amount reported.

Radium-equivalent uranium (RaeU), thorium (eTh), and potassium (eK) values (table 2) were determined by sealed-can gamma-ray spectrometry on 600 g of coarsely crushed (-32 mesh) sample (Bunker and Bush, 1966, 1967). The prefix "e" is used to designate this analytical method. The potassium value obtained by this technique is an actual measure of potassium and is generally precise and accurate (2σ) to within \pm the quantity [2 percent of the amount reported plus 0.03 percent absolute]. The large sample size used for gamma-ray spectrometry essentially eliminates splitting errors, and, therefore, eK is used throughout this report for calculation of radioelement ratios and statistical parameters.

Although eTh is not determined directly from thorium, disequilibrium within the thorium decay chain is unlikely. The precision and accuracy (2σ) for eTh is better than \pm the quantity (2 percent of the amount reported plus 0.1 ppm absolute). The value eTh is used in preference to Th for radioelement ratios and statistical calculations because of higher precision, lower detection limit, and smaller splitting error. RaeU is the amount of uranium needed for secular equilibrium with the indirectly measured ^{226}Ra . Radioactive

Table 1.—Chemical and normative compositions of granitic samples from the southeastern Arabian Shield [Chemical composition in percent. Letter symbol refers to location of fig. 1. D.I. is the differentiation index of Thornton and Tuttle, 1960. Fe₂O₃ and FeO calculated from the total iron assuming 2/3 of the iron is present as FeO. LOI is loss on ignition at 920°C]

| Symbol | 55500 A | 55501 A | 55502 A | 55503 A | 55504 A | 55505 A | 55506 B | 55507 B | 55508 B | 55509 B | 55510 B |
|--------------------------------|------------|------------|------------|------------|------------|------------|------------|------------|------------|------------|------------|
| SiO ₂ | 75.20 | 75.98 | 76.33 | 74.63 | 74.54 | 76.33 | 75.25 | 76.38 | 73.32 | 76.06 | 72.01 |
| Al ₂ O ₃ | 13.65 | 13.27 | 12.92 | 13.22 | 13.40 | 13.01 | 13.37 | 12.87 | 13.95 | 13.03 | 14.58 |
| Fe ₂ O ₃ | 0.19 | 0.21 | 0.19 | 0.49 | 0.47 | 0.24 | 0.37 | 0.32 | 0.47 | 0.38 | 0.58 |
| FeO | 0.34 | 0.38 | 0.34 | 0.89 | 0.84 | 0.43 | 0.67 | 0.58 | 0.84 | 0.69 | 1.05 |
| MgO | <0.10 | <0.10 | <0.10 | 0.25 | 0.26 | <0.10 | 0.18 | 0.17 | 0.34 | 0.16 | 0.52 |
| CaO | 0.30 | 0.16 | 0.25 | 0.88 | 0.92 | 0.24 | 0.60 | 0.50 | 1.09 | 0.51 | 1.52 |
| Na ₂ O | 4.47 | 4.39 | 3.92 | 3.54 | 3.69 | 4.09 | 4.16 | 3.86 | 4.29 | 3.92 | 4.31 |
| K ₂ O | 4.34 | 4.29 | 4.61 | 4.77 | 4.66 | 4.49 | 4.36 | 4.59 | 4.19 | 4.47 | 4.04 |
| LOI | 0.40 | 0.26 | 0.25 | 0.35 | 0.30 | 0.09 | 0.08 | 0.30 | 0.29 | 0.19 | 0.24 |
| TiO ₂ | <0.02 | 0.02 | <0.02 | 0.14 | 0.15 | <0.02 | 0.09 | 0.08 | 0.16 | 0.09 | 0.21 |
| P2O ₅ | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | 0.09 |
| MnO | 0.08 | 0.07 | 0.09 | 0.04 | 0.04 | 0.09 | 0.07 | 0.04 | 0.04 | 0.03 | 0.04 |
| ZrO ₂ | 0.01 | 0.01 | 0.01 | 0.02 | 0.02 | 0.01 | 0.01 | 0.01 | 0.02 | 0.01 | 0.02 |
| Total (-0) | 99.05 | 99.11 | 98.99 | 99.24 | 99.31 | 99.09 | 99.24 | 99.73 | 99.06 | 99.56 | 99.21 |
| Normative Minerals | | | | | | | | | | | |
| Q | 33.669 | 35.402 | 33.340 | 32.709 | 34.800 | 32.702 | 34.846 | 29.486 | 34.633 | 27.372 | |
| C | 1.112 | 1.172 | 1.086 | 0.687 | 0.666 | 1.043 | 0.770 | 0.693 | 0.524 | 0.868 | 0.573 |
| Z | 0.015 | 0.015 | 0.015 | 0.030 | 0.030 | 0.015 | 0.015 | 0.015 | 0.030 | 0.015 | 0.030 |
| OR | 25.892 | 25.579 | 27.519 | 28.403 | 27.730 | 26.776 | 25.963 | 27.198 | 24.996 | 26.530 | 24.063 |
| AB | 38.186 | 37.482 | 33.508 | 30.184 | 31.442 | 34.925 | 35.472 | 32.472 | 36.647 | 33.315 | 36.759 |
| AN | 1.371 | 0.669 | 1.121 | 4.267 | 4.464 | 1.070 | 2.868 | 2.356 | 5.063 | 2.410 | 7.008 |
| EN | 0.126 | 0.126 | 0.126 | 0.627 | 0.652 | 0.126 | 0.452 | 0.425 | 0.855 | 0.400 | 1.305 |
| FS | 0.601 | 0.649 | 0.629 | 1.075 | 0.991 | 0.745 | 0.915 | 0.747 | 0.977 | 0.862 | 1.184 |
| MT | 0.273 | 0.312 | 0.278 | 0.720 | 0.681 | 0.346 | 0.545 | 0.470 | 0.682 | 0.558 | 0.852 |
| IL | 0.019 | 0.019 | 0.019 | 0.268 | 0.287 | 0.019 | 0.172 | 0.152 | 0.307 | 0.172 | 0.402 |
| AP | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.143 | 0.048 | 0.215 |
| Total | 99.598 | 99.740 | 99.750 | 99.649 | 99.699 | 99.911 | 99.921 | 99.701 | 99.711 | 99.811 | 99.763 |
| SALIC | 98.531 | 98.586 | 98.650 | 96.910 | 97.041 | 98.628 | 97.790 | 96.746 | 97.771 | 95.805 | |
| FEMIC | 1.067 | 1.153 | 1.099 | 2.738 | 2.659 | 1.283 | 2.131 | 1.841 | 2.965 | 2.039 | 3.958 |
| D.I. | 96.034 | 96.730 | 96.428 | 91.926 | 91.881 | 96.501 | 94.137 | 94.796 | 91.129 | 94.479 | 88.194 |

Table 1.—Chemical and normative compositions of granitic samples from the southeastern Arabian Shield—Continued

| | 55511 | 55512 | 55513 | 55514 | 55515 | 55516 | 55517 | 55518 | 55519 | 55520 | 55521 |
|--------------------------------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|
| Symbol | C | C | C | C | C | C | C | D | D | D | D |
| SiO ₂ | 76.19 | 76.50 | 70.94 | 75.71 | 75.97 | 76.46 | 76.10 | 75.58 | 76.77 | 76.21 | 76.11 |
| Al ₂ O ₃ | 13.11 | 12.80 | 14.59 | 12.71 | 12.65 | 12.60 | 12.81 | 12.49 | 12.40 | 12.83 | 12.72 |
| Fe ₂ O ₃ | 0.28 | 0.30 | 0.77 | 0.37 | 0.39 | 0.32 | 0.33 | 0.57 | 0.59 | 0.32 | 0.61 |
| FeO | 0.51 | 0.54 | 1.39 | 0.67 | 0.71 | 0.57 | 0.59 | 1.03 | 0.58 | 0.58 | 0.74 |
| MgO | <0.10 | <0.10 | 0.46 | <0.10 | <0.10 | <0.10 | <0.10 | 0.15 | <0.10 | <0.10 | 0.05 |
| CaO | 0.28 | 0.38 | 0.73 | 0.69 | 0.54 | 0.48 | 0.60 | 0.39 | 0.39 | 0.39 | 0.41 |
| Na ₂ O | 4.35 | 4.22 | 3.44 | 3.90 | 3.87 | 3.76 | 3.97 | 3.05 | 3.92 | 4.25 | 4.14 |
| K ₂ O | 4.12 | 4.10 | 5.15 | 4.38 | 4.38 | 4.49 | 4.48 | 5.25 | 4.41 | 4.14 | 4.24 |
| Li ₂ O | 0.21 | 0.26 | 0.53 | 0.49 | 0.46 | 0.29 | 0.20 | 0.36 | 0.25 | 0.33 | 0.46 |
| TiO ₂ | <0.02 | <0.02 | 0.30 | 0.03 | 0.03 | <0.02 | 0.02 | 0.10 | 0.02 | 0.02 | 0.04 |
| P ₂ O ₅ | <0.05 | <0.05 | 0.13 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 |
| MnO | <0.02 | <0.02 | 0.03 | <0.02 | <0.02 | <0.02 | <0.02 | <0.02 | <0.02 | <0.02 | <0.02 |
| ZrO ₂ | 0.01 | 0.02 | 0.03 | 0.02 | 0.02 | 0.01 | 0.01 | 0.02 | 0.01 | 0.02 | 0.02 |
| Total (-0) | 99.15 | 99.21 | 99.24 | 99.09 | 99.25 | 99.13 | 99.06 | 99.23 | 99.15 | 99.17 | 99.37 |
| Normative Minerals | | | | | | | | | | | |
| Q | 34.469 | 35.361 | 27.256 | 34.583 | 35.023 | 36.184 | 34.775 | 35.987 | 36.188 | 34.717 | 34.674 |
| C | 1.042 | 0.783 | 0.985 | 0.277 | 0.339 | 0.626 | 0.611 | 0.753 | 0.521 | 0.702 | 0.627 |
| Z | 0.015 | 0.030 | 0.045 | 0.030 | 0.030 | 0.015 | 0.015 | 0.030 | 0.015 | 0.030 | 0.030 |
| OR | 24.554 | 24.421 | 30.667 | 26.121 | 26.078 | 26.766 | 26.724 | 31.266 | 26.284 | 24.670 | 25.215 |
| AB | 37.123 | 35.993 | 29.333 | 33.305 | 32.994 | 32.096 | 33.910 | 26.010 | 33.456 | 36.265 | 35.254 |
| AN | 1.269 | 1.768 | 6.543 | 3.523 | 3.317 | 2.571 | 2.272 | 2.868 | 1.820 | 1.819 | 1.915 |
| EN | 0.126 | 0.126 | 1.154 | 0.126 | 0.125 | 0.126 | 0.126 | 0.376 | 0.126 | 0.126 | 0.125 |
| FS | 0.711 | 0.752 | 1.482 | 0.895 | 0.952 | 0.795 | 0.804 | 1.277 | 0.786 | 0.786 | 0.976 |
| MT | 0.414 | 0.438 | 1.124 | 0.541 | 0.574 | 0.463 | 0.478 | 0.832 | 0.467 | 0.467 | 0.598 |
| IL | 0.019 | 0.019 | 0.574 | 0.058 | 0.057 | 0.019 | 0.038 | 0.191 | 0.038 | 0.038 | 0.076 |
| AP | 0.048 | 0.048 | 0.310 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 |
| Total | 99.789 | 99.739 | 99.473 | 99.507 | 99.538 | 99.709 | 99.799 | 99.638 | 99.749 | 99.668 | 99.538 |
| SALIC | 98.472 | 98.356 | 94.828 | 97.840 | 97.781 | 98.259 | 98.306 | 96.913 | 98.284 | 98.203 | 97.715 |
| FEMIC | 1.317 | 1.383 | 4.645 | 1.667 | 1.757 | 1.450 | 1.493 | 2.725 | 1.465 | 1.465 | 1.823 |
| D.I. | 96.146 | 95.775 | 87.256 | 94.010 | 94.095 | 95.047 | 95.408 | 93.262 | 95.928 | 95.652 | 95.143 |

Table 1.—Chemical and normative compositions of granitic samples from the southeastern Arabian Shield—Continued

| Symbol | 55522 E | 55523 E | 55524 E | 55525 E | 55526 F | 55527 F | 55528 F | 55529 F | 55530 G | 55531 G | 55532 H |
|--------------------------------|------------|------------|------------|------------|------------|------------|------------|------------|------------|------------|------------|
| SiO ₂ | 73.08 | 75.68 | 75.85 | 76.00 | 77.56 | 76.34 | 76.02 | 76.32 | 73.95 | 75.44 | 75.95 |
| Al ₂ O ₃ | 14.27 | 12.60 | 12.84 | 12.91 | 11.71 | 12.77 | 12.70 | 12.60 | 13.73 | 13.09 | 13.96 |
| Fe ₂ O ₃ | 0.29 | 0.47 | 0.32 | 0.33 | 0.50 | 0.37 | 0.48 | 0.49 | 0.52 | 0.44 | 0.32 |
| FeO | 0.53 | 0.85 | 0.58 | 0.60 | 0.89 | 0.66 | 0.87 | 0.83 | 0.94 | 0.80 | 0.58 |
| MgO | 0.15 | 0.11 | <0.10 | <0.10 | <0.10 | 0.11 | 0.13 | <0.10 | 0.29 | 0.20 | 0.05 |
| CaO | 0.64 | 0.70 | 0.66 | 0.61 | 0.34 | 0.52 | 0.56 | 0.49 | 1.33 | 0.86 | 0.26 |
| Na ₂ O | 3.24 | 3.09 | 3.63 | 3.41 | 3.30 | 3.67 | 3.66 | 3.63 | 4.26 | 3.90 | 3.98 |
| K ₂ O | 6.48 | 4.78 | 4.51 | 4.57 | 4.50 | 4.68 | 4.57 | 4.64 | 3.58 | 4.00 | 3.66 |
| Li ₂ O | 0.64 | 0.53 | 0.58 | 0.49 | 0.41 | 0.21 | 0.32 | 0.30 | 0.53 | 0.37 | 0.52 |
| TiO ₂ | 0.06 | 0.06 | 0.03 | 0.03 | 0.06 | 0.06 | 0.08 | 0.06 | 0.19 | 0.13 | <0.02 |
| P ₂ O ₅ | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 |
| MnO | <0.02 | 0.02 | 0.03 | 0.02 | 0.02 | 0.02 | 0.03 | 0.03 | 0.03 | <0.02 | 0.17 |
| ZrO ₂ | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.02 | 0.02 | 0.02 | 0.02 |
| Total (-0) | 99.42 | 98.92 | 99.11 | 99.05 | 99.37 | 99.44 | 99.45 | 99.52 | 99.40 | 99.49 | 99.49 |
| Normative Minerals | | | | | | | | | | | |
| Q | 27.758 | 37.708 | 35.996 | 37.329 | 40.112 | 35.671 | 35.583 | 36.013 | 31.955 | 35.195 | 37.911 |
| C | 0.815 | 1.130 | 0.862 | 1.305 | 0.846 | 0.774 | 0.767 | 0.767 | 0.480 | 0.835 | 3.042 |
| Z | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 | 0.030 | 0.030 | 0.030 | 0.030 |
| OR | 38.515 | 28.556 | 26.891 | 27.264 | 26.760 | 27.812 | 27.154 | 27.551 | 21.284 | 23.808 | 21.740 |
| AB | 27.576 | 26.433 | 30.993 | 29.130 | 28.101 | 31.231 | 31.140 | 30.863 | 36.266 | 33.240 | 33.852 |
| AN | 3.062 | 3.379 | 3.172 | 2.923 | 1.566 | 2.463 | 2.662 | 2.511 | 6.507 | 4.166 | 1.165 |
| EN | 0.376 | 0.277 | 0.126 | 0.126 | 0.125 | 0.276 | 0.326 | 0.125 | 0.502 | 0.502 | 0.125 |
| FS | 0.651 | 1.116 | 0.807 | 0.822 | 1.178 | 0.853 | 1.129 | 1.170 | 1.047 | 0.910 | 1.108 |
| MT | 0.427 | 0.688 | 0.468 | 0.487 | 0.724 | 0.534 | 0.704 | 0.703 | 0.763 | 0.647 | 0.471 |
| IL | 0.115 | 0.115 | 0.057 | 0.058 | 0.115 | 0.153 | 0.114 | 0.363 | 0.249 | 0.019 | |
| AP | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 | 0.048 |
| Total | 99.357 | 99.465 | 99.416 | 99.507 | 99.589 | 99.790 | 99.680 | 99.700 | 99.468 | 99.628 | 99.481 |
| SALIC | 97.741 | 97.221 | 97.910 | 97.966 | 97.399 | 97.966 | 97.321 | 97.535 | 96.521 | 97.273 | 97.710 |
| FEMIC | 1.616 | 2.245 | 1.506 | 1.541 | 2.189 | 1.824 | 2.359 | 2.165 | 2.947 | 2.355 | 1.770 |
| D.I. | 93.849 | 92.697 | 93.881 | 93.722 | 94.973 | 94.714 | 93.877 | 94.427 | 89.505 | 92.242 | 93.503 |

Table 1.—Chemical and normative compositions of granitic samples from the southeastern Arabian Shield—Continued

| Symbol | H | 55533 | 55534 | 55535 | 55536 | 55537 | 55538 | 55539 | 55540 | 55541 | 55542 | 55543 |
|--------------------------------|------------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|
| | | 1 | 1 | 1 | 1 | 1 | 1 | 1 | 1 | 1 | 1 | 1 |
| SiO ₂ | 76.43 | 75.72 | 76.47 | 78.19 | 80.19 | 76.99 | 76.40 | 69.25 | 69.80 | 74.08 | 66.66 | 66.66 |
| Al ₂ O ₃ | 13.73 | 13.94 | 13.84 | 11.44 | 10.89 | 11.90 | 12.57 | 13.85 | 13.78 | 13.04 | 14.81 | 14.81 |
| Fe ₂ O ₃ | 0.33 | 0.31 | 0.32 | 0.47 | 0.22 | 0.59 | 0.44 | 1.62 | 1.30 | 0.76 | 1.76 | 1.76 |
| FeO | 0.60 | 0.56 | 0.58 | 0.84 | 0.40 | 1.06 | 0.79 | 2.55 | 2.33 | 1.36 | 3.17 | 3.17 |
| MgO | <0.10 | <0.10 | <0.10 | <0.10 | <0.10 | <0.10 | 0.17 | 0.14 | <0.10 | 0.17 | 0.13 | 0.13 |
| CaO | 0.29 | 0.40 | 0.29 | 0.16 | 0.12 | 0.37 | 0.61 | 0.96 | 0.95 | 1.07 | 1.36 | 1.36 |
| Na ₂ O | 3.84 | 3.87 | 4.01 | 3.25 | 3.80 | 3.40 | 3.34 | 4.30 | 4.29 | 4.04 | 4.50 | 4.50 |
| K ₂ O | 3.48 | 3.88 | 3.10 | 4.69 | 3.58 | 4.88 | 4.90 | 5.59 | 5.45 | 4.09 | 5.67 | 5.67 |
| LOI | 0.69 | 0.69 | 0.91 | 0.33 | 0.28 | 0.27 | 0.27 | 0.24 | 0.18 | 0.18 | 0.26 | 0.26 |
| TiO ₂ | <0.02 | <0.02 | 0.02 | 0.06 | <0.02 | 0.07 | 0.12 | 0.36 | 0.31 | 0.10 | 0.48 | 0.48 |
| P ₂ O ₅ | <0.05 | 0.10 | 0.13 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | 0.05 | 0.05 |
| MnO | 0.17 | <0.02 | 0.04 | <0.02 | <0.02 | 0.02 | 0.02 | 0.02 | 0.07 | 0.06 | 0.04 | 0.09 |
| ZrO ₂ | Total (-0) | 99.64 | 99.54 | 99.75 | 99.53 | 99.60 | 99.64 | 99.64 | 98.87 | 98.66 | 99.08 | 99.08 |
| Normative Minerals | | | | | | | | | | | | |
| Q | 39.766 | 37.534 | 40.668 | 40.669 | 44.026 | 37.220 | 36.430 | 20.007 | 21.354 | 31.965 | 14.824 | 14.824 |
| C | 3.178 | 2.900 | 3.681 | 0.778 | 0.596 | 0.401 | 0.713 | 0.015 | 0.136 | 0.070 | 0.030 | 0.195 |
| Z | | | | | | | | | | | | |
| OR | 20.638 | 23.034 | 18.365 | 27.846 | 21.241 | 28.941 | 29.059 | 33.412 | 32.643 | 24.421 | 33.818 | 33.818 |
| AB | 32.609 | 32.899 | 34.018 | 27.631 | 32.285 | 28.873 | 28.364 | 36.803 | 36.794 | 34.562 | 38.433 | 38.433 |
| AN | 1.313 | 1.337 | 0.591 | 0.666 | 0.467 | 1.711 | 2.906 | 2.003 | 2.278 | 5.232 | 3.498 | 3.498 |
| WO | | | | | | | | | | | | |
| EN | 0.125 | 0.125 | 0.125 | 0.125 | 0.125 | 0.125 | 0.425 | 0.353 | 0.227 | 0.428 | 0.327 | 0.327 |
| FS | 1.131 | 0.775 | 0.854 | 1.082 | 0.558 | 1.390 | 0.907 | 3.085 | 2.855 | 1.805 | 3.784 | 3.784 |
| MT | 0.485 | 0.451 | 0.465 | 0.679 | 0.325 | 0.858 | 0.635 | 2.076 | 1.904 | 1.107 | 2.578 | 2.578 |
| IL | 0.019 | 0.019 | 0.019 | 0.114 | 0.019 | 0.133 | 0.229 | 0.692 | 0.597 | 0.192 | 0.920 | 0.920 |
| AP | 0.048 | 0.238 | 0.309 | 0.048 | 0.048 | 0.048 | 0.048 | 0.072 | 0.072 | 0.068 | 0.120 | 0.120 |
| Total | 99.311 | 99.095 | 99.670 | 99.720 | 99.730 | 99.730 | 99.760 | 99.820 | 99.820 | 99.741 | 99.741 | 99.741 |
| SALIC | 97.504 | 97.324 | 97.621 | 98.645 | 97.176 | 97.487 | 92.390 | 93.205 | 96.240 | 90.767 | 8.974 | 8.974 |
| FEMIC | 1.807 | 1.608 | 1.771 | 2.049 | 1.075 | 2.554 | 2.243 | 7.369 | 6.615 | 3.580 | | |
| D.I. | 93.013 | 93.467 | 93.052 | 96.147 | 97.553 | 95.034 | 93.853 | 90.222 | 90.791 | 90.908 | 87.074 | 87.074 |

Table 1.—Chemical and normative compositions of granitic samples from the southeastern Arabian Shield—Continued

| Symbol | 55544 L | 55545 L | 55546 M | 55547 M | 55548 M | 55549 M | 55550 N | 55551 P | 55552 P | 55553 Q | 55554 Q |
|--------------------|------------|------------|------------|------------|------------|------------|------------|------------|------------|------------|------------|
| S102 | 71.86 | 71.23 | 75.41 | 74.17 | 74.62 | 72.80 | 74.40 | 75.51 | 75.74 | 75.30 | 78.84 |
| Al2O3 | 13.57 | 13.79 | 13.26 | 13.98 | 14.19 | 15.60 | 14.16 | 13.04 | 12.63 | 13.99 | 11.50 |
| Fe2O3 | 0.35 | 0.95 | 0.93 | 0.16 | 0.16 | 0.12 | 0.33 | 0.39 | 0.49 | 0.19 | 0.44 |
| Fe | 1.53 | 1.71 | 1.67 | 0.28 | 0.29 | 0.22 | 0.60 | 0.70 | 0.89 | 0.35 | 0.79 |
| MgO | 0.28 | 0.28 | <0.10 | <0.10 | <0.10 | <0.10 | 0.12 | <0.10 | <0.10 | <0.10 | <0.10 |
| CaO | 1.02 | 1.05 | 0.28 | 0.70 | 0.34 | 0.35 | 0.67 | 0.54 | 0.61 | 0.12 | 0.17 |
| Na2O | 3.88 | 3.94 | 4.31 | 4.82 | 4.86 | 4.93 | 4.19 | 4.28 | 3.70 | 5.14 | 3.69 |
| K2O | 4.87 | 4.88 | 4.21 | 3.84 | 3.73 | 4.23 | 4.14 | 4.31 | 4.44 | 4.00 | 3.86 |
| LOI | 0.71 | 0.57 | 0.41 | 0.53 | 0.44 | 0.45 | 0.45 | 0.50 | 0.44 | 0.09 | 0.22 |
| TiO2 | 0.20 | 0.23 | <0.02 | <0.02 | <0.02 | <0.02 | <0.02 | 0.03 | 0.03 | <0.02 | <0.02 |
| P2O5 | 0.07 | 0.08 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 |
| MnO | 0.04 | 0.05 | 0.04 | 0.04 | 0.05 | 0.08 | 0.07 | <0.02 | <0.02 | <0.02 | <0.02 |
| ZrO2 | 0.03 | 0.03 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.02 | 0.02 | 0.01 | 0.02 |
| Total (-0) | 98.91 | 98.79 | 100.59 | 98.61 | 98.78 | 98.86 | 99.23 | 99.38 | 99.11 | 99.28 | 99.64 |
| Normative Minerals | | | | | | | | | | | |
| Q | 27.692 | 26.560 | 32.314 | 30.131 | 31.491 | 27.284 | 32.571 | 32.720 | 35.696 | 29.810 | 41.847 |
| C | 0.232 | 0.313 | 1.121 | 0.679 | 1.607 | 2.350 | 1.725 | 0.403 | 0.682 | 1.042 | 0.994 |
| Z | 0.045 | 0.045 | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 | 0.030 | 0.015 | 0.030 |
| OR | 29.095 | 29.191 | 24.731 | 23.012 | 22.314 | 25.286 | 24.653 | 25.627 | 26.472 | 23.808 | 22.892 |
| AB | 33.193 | 33.748 | 36.255 | 41.361 | 41.633 | 42.199 | 35.729 | 36.441 | 31.589 | 43.808 | 31.336 |
| AN | 4.654 | 4.744 | 1.316 | 3.389 | 1.575 | 1.624 | 2.955 | 2.564 | 2.922 | 0.468 | 0.715 |
| EN | 0.705 | 0.706 | 0.124 | 0.126 | 0.126 | 0.126 | 0.301 | 0.125 | 0.126 | 0.125 | 0.125 |
| FS | 1.873 | 2.095 | 2.351 | 0.453 | 0.488 | 0.435 | 0.915 | 0.934 | 1.137 | 0.485 | 1.135 |
| MT | 1.245 | 1.393 | 1.339 | 0.230 | 0.240 | 0.176 | 0.487 | 0.564 | 0.721 | 0.282 | 0.640 |
| IL | 0.384 | 0.442 | 0.019 | 0.019 | 0.019 | 0.019 | 0.057 | 0.057 | 0.134 | 0.019 | 0.019 |
| AP | 0.168 | 0.192 | 0.024 | 0.048 | 0.048 | 0.048 | 0.143 | 0.048 | 0.048 | 0.048 | 0.048 |
| Total | 99.286 | 99.428 | 99.593 | 99.464 | 99.556 | 99.547 | 99.551 | 99.498 | 99.557 | 99.910 | 99.781 |
| SALIC | 94.912 | 94.600 | 95.737 | 98.588 | 98.636 | 98.743 | 97.648 | 97.770 | 97.392 | 98.951 | 97.814 |
| FEMIC | 4.375 | 4.828 | 3.857 | 0.877 | 0.920 | 0.804 | 1.903 | 1.728 | 2.166 | 0.960 | 1.966 |
| D.I. | 89.581 | 89.498 | 93.300 | 94.504 | 95.438 | 94.769 | 92.953 | 94.788 | 93.758 | 97.426 | 96.075 |

Table 1.—Chemical and normative compositions of granitic samples from the southeastern Arabian Shield—Continued

| | 55555 | 55556 | 55557 | 55558 | 55559 | 55560 | 55561 | 55562 | 55563 | 55564 | 55565 |
|--------------------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|--------|
| Symbol | Q | Q | R | R | R | S | T | T | U | U | U |
| Si02 | 76.14 | 73.14 | 71.29 | 67.76 | 74.45 | 72.39 | 73.31 | 74.09 | 70.35 | 75.44 | 75.05 |
| Al2O3 | 13.38 | 13.89 | 16.03 | 17.24 | 14.84 | 15.68 | 15.27 | 14.78 | 15.80 | 13.89 | 14.15 |
| Fe2O3 | 0.13 | 0.55 | 0.53 | 0.83 | 0.17 | 0.30 | 0.13 | 0.21 | 0.68 | 0.17 | 0.25 |
| FeO | 0.23 | 0.98 | 0.96 | 1.49 | 0.31 | 0.55 | 0.23 | 0.37 | 1.23 | 0.30 | 0.44 |
| MgO | <0.10 | 0.33 | 0.58 | 1.14 | <0.10 | 0.28 | <0.10 | 0.25 | 0.83 | <0.10 | <0.10 |
| CaO | 0.17 | 0.83 | 3.52 | 4.58 | 1.28 | 2.50 | 1.23 | 1.00 | 2.68 | 0.55 | 0.53 |
| Na2O | 4.85 | 3.54 | 4.72 | 4.63 | 4.28 | 4.76 | 4.03 | 4.34 | 4.39 | 4.17 | 4.51 |
| K2O | 3.55 | 4.79 | 0.90 | 1.04 | 3.66 | 2.07 | 4.44 | 3.14 | 2.39 | 4.26 | 4.05 |
| LOI | 0.29 | 0.91 | 0.57 | 0.30 | 0.30 | 0.36 | 0.13 | 0.66 | 0.45 | 0.25 | 0.24 |
| TiO2 | <0.02 | 0.19 | 0.17 | 0.28 | 0.02 | 0.09 | <0.02 | 0.03 | <0.02 | <0.02 | <0.02 |
| P2O5 | <0.05 | 0.05 | 0.08 | 0.13 | <0.05 | <0.05 | <0.05 | 0.09 | 0.11 | <0.05 | <0.05 |
| MnO | 0.05 | <0.02 | 0.02 | 0.04 | 0.03 | <0.02 | 0.12 | 0.09 | 0.06 | 0.08 | 0.06 |
| ZrO2 | 0.01 | 0.02 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 |
| Total (-0) | 98.88 | 99.23 | 99.37 | 99.46 | 99.41 | 99.01 | 98.97 | 99.05 | 99.23 | 99.20 | 99.37 |
| Normative Minerals | | | | | | | | | | | |
| Q | 34.156 | 31.841 | 31.922 | 25.075 | 32.769 | 31.075 | 30.293 | 34.610 | 28.579 | 33.703 | 32.039 |
| C | 1.312 | 1.504 | 1.090 | 0.484 | 1.568 | 1.123 | 1.663 | 2.664 | 1.393 | 1.479 | 1.440 |
| Z | 0.015 | 0.030 | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 | 0.015 |
| OR | 21.215 | 28.525 | 5.352 | 6.179 | 21.757 | 12.355 | 26.509 | 18.733 | 14.232 | 25.377 | 24.084 |
| AB | 41.503 | 30.187 | 40.191 | 39.389 | 36.433 | 40.681 | 34.454 | 37.077 | 37.434 | 35.571 | 38.404 |
| AN | 0.721 | 3.820 | 17.047 | 6.257 | 12.395 | 6.033 | 4.415 | 12.674 | 2.619 | 2.514 | |
| EN | 0.126 | 0.828 | 1.454 | 2.854 | 0.125 | 0.704 | 0.126 | 0.629 | 2.083 | 0.126 | 0.125 |
| FS | 0.404 | 1.069 | 1.068 | 1.672 | 0.472 | 0.645 | 0.535 | 0.637 | 1.405 | 0.550 | 0.712 |
| MT | 0.190 | 0.798 | 0.777 | 1.204 | 0.253 | 0.444 | 0.190 | 0.302 | 0.997 | 0.243 | 0.360 |
| IL | 0.019 | 0.364 | 0.325 | 0.535 | 0.019 | 0.153 | 0.019 | 0.058 | 0.478 | 0.019 | 0.019 |
| AP | 0.048 | 0.119 | 0.191 | 0.310 | 0.048 | 0.048 | 0.048 | 0.215 | 0.263 | 0.048 | 0.048 |
| Total | 99.708 | 99.086 | 99.431 | 99.706 | 99.700 | 99.638 | 99.871 | 99.553 | 99.750 | 99.760 | |
| SALIC | 98.921 | 95.907 | 95.616 | 93.131 | 98.783 | 97.643 | 98.953 | 97.499 | 94.327 | 98.764 | 98.497 |
| FEMIC | 0.787 | 3.179 | 3.814 | 6.574 | 0.917 | 1.995 | 0.918 | 1.840 | 5.226 | 0.986 | 1.263 |
| D.I. | 96.873 | 90.553 | 77.465 | 70.642 | 90.959 | 84.111 | 91.257 | 90.420 | 80.245 | 94.651 | 94.527 |

Table 1.—Chemical and normative compositions of granitic samples from the southeastern Arabian Shield--Continued

| Symbol | 55566 U | 55567 V | 55568 V | 55569 V | 55570 V | 55571 V | 55572 V |
|--------------------------------|------------|------------|------------|------------|------------|------------|------------|
| SiO ₂ | 76.69 | 75.00 | 75.19 | 75.21 | 75.02 | 74.98 | 74.52 |
| Al ₂ O ₃ | 12.99 | 11.87 | 11.99 | 12.49 | 13.64 | 11.95 | 13.56 |
| Fe ₂ O ₃ | 0.20 | 0.81 | 0.84 | 0.57 | 0.28 | 0.97 | 0.28 |
| FeO | 0.36 | 1.46 | 1.52 | 1.02 | 0.50 | 1.74 | 0.50 |
| MgO | <0.10 | <0.10 | <0.10 | 0.11 | <0.10 | 0.12 | 0.12 |
| CaO | 0.54 | 0.53 | 0.52 | 0.53 | 0.72 | 0.54 | 0.33 |
| Na ₂ O | 3.83 | 3.55 | 3.41 | 3.72 | 3.71 | 3.42 | 3.88 |
| K ₂ O | 4.20 | 5.02 | 5.12 | 4.77 | 4.63 | 5.17 | 5.49 |
| Li ₂ O | 0.14 | 0.31 | 0.33 | 0.35 | 0.53 | 0.19 | 0.38 |
| TiO ₂ | <0.02 | 0.20 | 0.21 | 0.12 | 0.06 | 0.25 | 0.05 |
| P ₂ O ₅ | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 | <0.05 |
| MnO | 0.04 | 0.06 | 0.06 | 0.03 | 0.04 | 0.07 | <0.02 |
| ZrO ₂ | | 0.07 | 0.07 | 0.05 | 0.01 | 0.08 | 0.02 |
| Total (-0) | 99.08 | 93.96 | 99.33 | 98.98 | 99.20 | 99.50 | 99.15 |
| Normative Minerals | | | | | | | |
| Q | 37.196 | 33.672 | 34.011 | 33.792 | 34.126 | 33.309 | 30.073 |
| C | 1.197 | | | 0.271 | 1.274 | | 0.689 |
| Z | | 0.105 | 0.105 | 0.075 | 0.015 | 0.120 | 0.030 |
| OR | 25.049 | 29.977 | 30.459 | 28.479 | 27.579 | 30.706 | 32.719 |
| AB | 32.709 | 30.356 | 29.049 | 31.803 | 31.645 | 29.086 | 33.111 |
| AN | 2.638 | 1.645 | 2.304 | 2.591 | 3.469 | 1.997 | 1.519 |
| WO | | 0.368 | 0.068 | | | 0.236 | |
| EN | 0.176 | 0.126 | 0.125 | 0.277 | 0.126 | 0.300 | 0.301 |
| FS | 0.559 | 1.813 | 1.870 | 1.277 | 0.667 | 2.126 | 0.628 |
| MT | 0.292 | 1.190 | 1.230 | 0.829 | 0.404 | 1.407 | 0.404 |
| IL | 0.019 | 0.384 | 0.402 | 0.230 | 0.115 | 0.477 | 0.096 |
| AP | 0.024 | 0.048 | 0.048 | 0.024 | 0.048 | 0.048 | 0.048 |
| Total | 99.360 | 99.689 | 99.670 | 99.647 | 99.467 | 99.811 | 99.618 |
| SALIC | 98.789 | 95.755 | 95.928 | 97.010 | 98.108 | 95.216 | 98.141 |
| FEMIC | 1.071 | 3.934 | 3.742 | 2.637 | 1.359 | 4.595 | 1.477 |
| D.I. | 94.955 | 94.005 | 93.520 | 94.074 | 93.356 | 93.100 | 95.903 |

Table 2.—Trace-element concentrations in granite samples from the southeastern Arabian Shield
 [Letters in parentheses after pluton names refer to locations on fig. 1. Abbreviations are defined
 in text]

| Sample | Latitude (north) | Longitude (east) | U (ppm) | Ra ^{eU} (ppm) | Th (ppm) | e^{Th} (ppm) | K (wt %) | e^{K} (wt %) | Cu (ppm) | Zn (ppm) | Rb (ppm) | Sr (ppm) |
|--------|------------------|------------------|---------|------------------------|------------------------------|------------------------------|----------|------------------------------|----------|----------|----------|----------|
| 155500 | 21 15 30 | 43 58 15 | 7.19 | 6.1 | 34.20 | 28.2 | 3.60 | 3.66 | 7.0 | 171 | 740 | 9 |
| 155501 | 21 15 30 | 43 58 0 | 10.70 | 8.9 | 28.40 | 28.9 | 3.56 | 3.62 | 7.8 | 164 | 759 | 6 |
| 155502 | 21 15 30 | 43 58 30 | 8.84 | 6.7 | 28.40 | 28.3 | 3.83 | 3.84 | 6.2 | 168 | 824 | 7 |
| 155503 | 21 14 40 | 44 0 0 | 8.04 | 7.9 | 24.40 | 22.7 | 3.96 | 3.99 | 5.6 | 95 | 303 | 93 |
| 155504 | 21 14 40 | 44 0 15 | 10.10 | 7.5 | 27.10 | 24.3 | 3.87 | 3.95 | 3.4 | 83 | 265 | 95 |
| 155505 | 21 13 45 | 44 2 15 | 7.76 | 6.7 | 27.80 | 27.3 | 3.73 | 3.72 | 6.3 | 166 | 780 | 10 |
| | | | | | Jabal Tarban (A) | | | | | | | |
| 155506 | 21 42 0 | 43 43 15 | 3.22 | 2.6 | 21.20 | 18.8 | 3.62 | 3.66 | 4.9 | 4.8 | 121 | 65 |
| 155507 | 21 42 15 | 43 44 0 | 1.76 | 2.0 | 8.80 | 10.5 | 3.81 | 3.73 | 6.5 | 6.9 | 102 | 52 |
| 155508 | 21 43 30 | 43 42 45 | 4.24 | 4.7 | 14.00 | 11.6 | 3.48 | 3.50 | 6.4 | 7.8 | 131 | 205 |
| 155509 | 21 43 30 | 43 41 15 | 2.23 | 2.3 | 9.71 | 10.0 | 3.71 | 3.74 | 4.7 | 6.5 | 102 | 72 |
| 155510 | 21 43 20 | 43 39 20 | 3.64 | 3.6 | 13.50 | 14.9 | 3.35 | 3.34 | 2.8 | 5.5 | 125 | 371 |
| | | | | | Jabal as Sukkah (B) | | | | | | | |
| 155511 | 23 17 40 | 44 35 55 | 8.18 | 7.7 | 28.40 | 29.4 | 3.42 | 3.52 | 32 | 168 | 685 | 3 |
| 155512 | 23 17 45 | 44 35 55 | 13.80 | 12.5 | 46.20 | 44.4 | 3.40 | 3.41 | 72 | 71 | 490 | 3 |
| 155513 | 23 17 50 | 44 35 10 | 4.78 | 4.6 | 18.90 | 18.5 | 4.28 | 4.31 | 33 | 63 | 182 | 187 |
| 155514 | 23 17 50 | 44 37 45 | 18.20 | 16.0 | 53.40 | 49.1 | 3.64 | 3.76 | 39 | 102 | 390 | 3 |
| 155515 | 23 17 10 | 44 37 45 | 20.00 | 18.0 | 56.10 | 53.5 | 3.64 | 3.69 | 85 | 77 | 337 | 8 |
| 155516 | 23 15 10 | 44 38 30 | 8.40 | 7.1 | 45.20 | 44.2 | 3.73 | 3.85 | 31 | 69 | 425 | 4 |
| 155517 | 23 15 10 | 44 38 30 | 9.27 | 8.2 | 43.70 | 41.0 | 3.72 | 3.88 | 4.8 | 145 | 430 | 3 |
| | | | | | Jabal Sabbah (C) | | | | | | | |
| 155518 | 22 53 30 | 44 51 58 | 7.18 | 7.1 | 21.90 | 27.7 | 4.36 | 4.37 | 6.1 | 83 | 211 | 29 |
| 155519 | 22 49 50 | 44 51 28 | 4.16 | 4.8 | 16.30 | 18.7 | 3.66 | 3.77 | 3.1 | 49 | 246 | <1 |
| 155520 | 22 49 55 | 44 51 30 | 9.32 | 9.2 | 36.40 | 33.6 | 3.44 | 3.56 | 8.6 | 164 | 610 | 5 |
| 155521 | 22 46 15 | 44 51 25 | 9.66 | 8.8 | 23.50 | 23.4 | 3.52 | 3.60 | 71 | 104 | 297 | 1 |
| | | | | | Jabal Hawshat ibn Hawayl (D) | | | | | | | |
| 155522 | 22 38 45 | 44 47 10 | 9.53 | 7.1 | 23.40 | 20.7 | 5.38 | 5.41 | 4.4 | 34 | 441 | 73 |
| 155523 | 22 37 30 | 44 43 15 | 15.20 | 13.4 | 46.00 | 44.1 | 3.97 | 3.98 | 79 | 79 | 592 | 16 |
| 155524 | 22 31 58 | 44 45 45 | 25.10 | 19.0 | 29.30 | 31.6 | 3.74 | 3.79 | 29 | 63 | 609 | 8 |
| 155525 | 22 31 55 | 44 45 55 | 15.90 | 15.9 | 43.30 | 37.8 | 3.79 | 3.89 | 64 | 76 | 532 | 14 |
| | | | | | Jabal al Hawshaw (E) | | | | | | | |
| 155526 | 22 12 10 | 44 48 15 | 6.05 | 8.4 | 17.60 | 26.7 | 3.74 | 3.82 | 3.8 | 145 | 372 | 1 |
| 155527 | 22 11 0 | 44 49 15 | 9.33 | 9.5 | 20.30 | 21.8 | 3.89 | 4.00 | 24 | 84 | 212 | 20 |
| 155528 | 22 9 5 | 44 49 30 | 7.25 | 7.8 | 21.50 | 23.7 | 3.79 | 4.00 | 71 | 91 | 257 | 41 |
| 155529 | 22 6 40 | 44 52 0 | 9.28 | 9.0 | 21.10 | 25.8 | 3.85 | 3.97 | 63 | 118 | 317 | 14 |
| | | | | | Jabal Sahah (F) | | | | | | | |
| 155530 | 22 6 0 | 44 43 35 | 2.18 | 2.4 | 8.51 | 10.4 | 2.97 | 3.10 | 78 | 64 | 99 | 100 |
| 155531 | 22 6 0 | 44 43 35 | 3.08 | 3.02 | 13.80 | 12.7 | 3.32 | 3.44 | 73 | 63 | 112 | 64 |

Table 2.—Trace-element concentrations in granite samples from the southeastern Arabian Shield--Continued

| Sample | Y (ppm) | Zr (ppm) | Th/U | eTh/U | K/U | eK/U | Th/K | eTh/eK | K/Rb | Rb/Sr | Al/(Na+K) | Al/(Na+K+Ca) |
|------------------------------|------------|-------------|------|-------|------|------|-------|--------|--------|---------|-----------|--------------|
| 155500 | 87 | 72 | 4.76 | 3.92 | .50 | .51 | 9.49 | 7.70 | 48.69 | 82.22 | 1.13 | 1.08 |
| 155501 | 102 | 80 | 2.65 | 2.70 | .33 | .34 | 7.97 | 7.98 | 46.92 | 126.50 | 1.12 | 1.09 |
| 155502 | 81 | 70 | 3.21 | 3.20 | .43 | .43 | 7.42 | 7.37 | 46.44 | 117.71 | 1.13 | 1.09 |
| 155503 | 36 | 121 | 3.03 | 2.82 | .49 | .50 | 6.16 | 5.69 | 130.69 | 3.26 | 1.20 | 1.05 |
| 155504 | 47 | 125 | 2.68 | 2.41 | .38 | .39 | 7.01 | 6.15 | 145.98 | 2.79 | 1.21 | 1.05 |
| 155505 | 78 | 68 | 3.58 | 3.52 | .48 | .48 | 7.46 | 7.34 | 47.79 | 78.00 | 1.12 | 1.08 |
| Jabal as Sukkah (B) | | | | | | | | | | | | |
| 155506 | 8 | 96 | 6.58 | 5.84 | 1.12 | 1.14 | 5.86 | 5.14 | 299.13 | 1.86 | 1.16 | 1.06 |
| 155507 | 6 | 91 | 5.00 | 5.97 | 2.17 | 2.12 | 2.31 | 2.82 | 373.57 | 1.96 | 1.14 | 1.05 |
| 155508 | 19 | 139 | 3.30 | 2.74 | .82 | .83 | 4.02 | 3.31 | 265.52 | -64 | 1.20 | 1.03 |
| 155509 | 7 | 83 | 4.35 | 4.48 | 1.66 | 1.68 | 2.62 | 2.67 | 363.81 | 1.42 | 1.15 | 1.07 |
| 155510 | 12 | 118 | 3.71 | 4.09 | .92 | .92 | 4.03 | 4.46 | 268.31 | .34 | 1.27 | 1.02 |
| Jabal Sabbah (C) | | | | | | | | | | | | |
| 155511 | 213 | 76 | 3.47 | 3.59 | .42 | .43 | 8.30 | 8.35 | 49.93 | 228.33 | 1.13 | 1.08 |
| 155512 | 221 | 112 | 3.35 | 3.22 | .25 | .25 | 13.57 | 13.02 | 69.46 | 163.33 | 1.12 | 1.06 |
| 155513 | 21 | 221 | 3.95 | 3.87 | .89 | .90 | 4.42 | 4.29 | 234.91 | .97 | 1.30 | 1.05 |
| 155514 | 170 | 115 | 2.93 | 2.70 | .20 | .21 | 14.69 | 13.06 | 93.23 | 130.00 | 1.14 | 1.02 |
| 155515 | 154 | 127 | 2.81 | 2.68 | .18 | .18 | 15.43 | 14.50 | 107.90 | 42.13 | 1.14 | 1.02 |
| 155516 | 137 | 103 | 5.38 | 5.26 | .44 | .46 | 12.13 | 11.48 | 87.70 | 106.25 | 1.14 | 1.05 |
| 155517 | 179 | 89 | 4.71 | 4.42 | .40 | .42 | 11.75 | 10.57 | 86.49 | 143.33 | 1.13 | 1.05 |
| Jabal Hawshat ibn Hawayl (D) | | | | | | | | | | | | |
| 155518 | 99 | 156 | 3.05 | 3.86 | .61 | .61 | 5.02 | 6.34 | 206.56 | 7.28 | 1.17 | 1.06 |
| 155519 | 87 | 83 | 3.92 | 4.50 | .88 | .91 | 4.45 | 4.96 | 148.82 | >246.00 | 1.10 | 1.04 |
| 155520 | 123 | 133 | 3.91 | 3.61 | .37 | .38 | 10.59 | 9.44 | 56.34 | 122.00 | 1.12 | 1.05 |
| 155521 | 133 | 143 | 2.43 | 2.42 | .36 | .37 | 6.68 | 6.50 | 118.51 | 297.00 | 1.12 | 1.05 |
| Jabal al Hawshaw (E) | | | | | | | | | | | | |
| 155522 | 32 | 69 | 2.46 | 2.17 | .56 | .57 | 4.35 | 3.83 | 121.98 | 6.04 | 1.16 | 1.06 |
| 155523 | 77 | 105 | 3.03 | 2.90 | .26 | .26 | 11.59 | 11.08 | 67.03 | 37.00 | 1.23 | 1.09 |
| 155524 | 107 | 60 | 1.12 | 1.21 | .14 | .15 | 7.83 | 8.34 | 61.48 | 76.13 | 1.18 | 1.07 |
| 155525 | 92 | 73 | 2.72 | 2.38 | .24 | .24 | 11.41 | 9.72 | 71.31 | 38.00 | 1.22 | 1.11 |
| Jabal Sahah (F) | | | | | | | | | | | | |
| 155526 | 77 | 79 | 2.91 | 4.41 | .62 | .63 | 4.71 | 6.99 | 100.42 | 372.00 | 1.14 | 1.07 |
| 155527 | 78 | 109 | 2.18 | 2.34 | .42 | .43 | 5.23 | 5.45 | 183.26 | 10.60 | 1.15 | 1.06 |
| 155528 | 68 | 95 | 2.97 | 3.27 | .52 | .55 | 5.67 | 5.93 | 147.62 | 6.27 | 1.16 | 1.06 |
| 155529 | 119 | 119 | 2.27 | 2.78 | .42 | .43 | 5.48 | 6.50 | 121.51 | 22.64 | 1.15 | 1.06 |
| Jabal as Sitarah (G) | | | | | | | | | | | | |
| 155530 | 22 | 112 | 3.90 | 4.77 | 1.36 | 1.42 | 2.86 | 3.35 | 300.20 | .99 | 1.26 | 1.03 |
| 155531 | 30 | 121 | 4.48 | 4.12 | 1.08 | 1.12 | 4.16 | 3.69 | 296.49 | 1.75 | 1.22 | 1.06 |

Table 2.—Trace-element concentrations in granite samples from the southeastern Arabian Shield—Continued

| Sample | Latitude (north) | Longitude (east) | U (ppm) | Ra <u>E</u> (ppm) | Th (ppm) | e <u>Th</u> (wt %) | K (wt %) | e <u>K</u> (wt %) | Cu (ppm) | Zn (ppm) | Rb (ppm) | $\frac{L_{Sr}}{L_{Ra}}$ (ppm) |
|--------|---------------------|---------------------|------------|----------------------|------------------------------|-----------------------|-------------|----------------------|-------------|-------------|-------------|----------------------------------|
| | | | | | Jabal Mahail (H) | | | | | | | |
| 155532 | 21 19 15 | 44 47 50 | 4.88 | 6.4 | 7.60 | 10.5 | 3.04 | 3.27 | 4.8 | 76 | 311 | 4 |
| 155533 | 21 19 15 | 44 47 50 | 2.56 | 2.6 | 7.52 | 10.8 | 2.89 | 3.23 | 6.2 | 75 | 337 | 7 |
| | | | | | Jabal Kebad (I) | | | | | | | |
| 155534 | 21 10 20 | 44 34 0 | 5.12 | 5.2 | 8.16 | 9.2 | 3.22 | 3.41 | 10.1 | 86 | 335 | 6 |
| 155535 | 21 12 35 | 44 31 30 | 2.43 | 3.1 | 4.60 | 8.1 | 2.57 | 2.82 | 137 | 122 | 492 | 8 |
| | | | | | Hudban (J) | | | | | | | |
| 155536 | 21 46 35 | 43 49 30 | 8.94 | 10.9 | 18.40 | 36.6 | 3.89 | 3.94 | 72 | 139 | 323 | 9 |
| 155537 | 21 46 35 | 43 49 30 | 26.90 | 31.0 | 73.70 | 106.0 | 2.97 | 2.91 | 57 | 84 | 291 | 5 |
| 155538 | 21 49 55 | 43 50 5 | 7.77 | 11.3 | 19.20 | 31.0 | 4.05 | 4.09 | 85 | 90 | 228 | 10 |
| 155539 | 21 48 20 | 43 53 45 | 7.56 | 10.8 | 18.60 | 31.7 | 4.07 | 4.05 | 77 | 90 | 321 | 69 |
| | | | | | Najran (K) | | | | | | | |
| 155540 | 17 33 30 | 44 0 0 | 4.14 | 4.0 | 12.40 | 11.2 | 4.64 | 4.69 | 67 | 147 | 127 | 70 |
| 155541 | 17 33 30 | 44 0 0 | 4.94 | 5.0 | 15.60 | 11.3 | 4.52 | 4.62 | 55 | 111 | 123 | 64 |
| 155542 | 17 32 45 | 44 9 15 | 4.58 | 3.4 | 16.50 | 15.5 | 3.40 | 3.52 | 53 | 70 | 100 | 87 |
| 155543 | 17 32 45 | 44 9 15 | 1.29 | 1.5 | 9.13 | 7.0 | 4.71 | 4.82 | 148 | 166 | 78 | 116 |
| | | | | | Jabal Ashirah (L) | | | | | | | |
| 155544 | 18 0 45 | 44 12 15 | 3.36 | 3.8 | 14.00 | 14.1 | 4.04 | 4.15 | 22 | 91 | 175 | 92 |
| 155545 | 18 0 45 | 44 12 15 | 4.46 | 4.2 | 19.60 | 14.3 | 4.05 | 4.06 | 60 | 82 | 181 | 106 |
| | | | | | Jabal al Gaharra (M) | | | | | | | |
| 155546 | 18 3 15 | 44 0 15 | 3.94 | 8.8 | 18.30 | 13.6 | 3.49 | 3.61 | 54 | 120 | 606 | 6 |
| 155547 | 18 3 30 | 44 0 30 | 15.80 | 14.3 | 21.40 | 18.2 | 3.19 | 3.19 | 44 | 182 | 651 | 3 |
| 155548 | 18 3 45 | 44 0 0 | 12.00 | 12.0 | 24.40 | 19.2 | 3.10 | 3.14 | 46 | 80 | 680 | 4 |
| 155549 | 18 5 30 | 44 0 30 | 2.63 | 1.9 | 19.00 | 13.5 | 3.51 | 3.58 | 74 | 127 | 952 | 4 |
| | | | | | Northeast of As Subaykah (N) | | | | | | | |
| 155550 | 19 10 40 | 43 23 15 | 4.61 | 4.5 | <3.20 | 3.1 | 3.44 | 3.53 | 87 | 83 | 84 | 130 |
| | | | | | East of As Subaykah (P) | | | | | | | |
| 155551 | 19 5 20 | 43 24 25 | 5.16 | 4.3 | 9.63 | 7.3 | 3.58 | 3.79 | 69 | 100 | 88 | 8 |
| 155552 | 19 7 10 | 43 23 25 | 2.92 | 2.2 | 6.00 | 7.8 | 3.69 | 3.93 | 65 | 90 | 77 | 40 |
| | | | | | Jabal Bani Bwana (Q) | | | | | | | |
| 155553 | 18 31 50 | 43 56 10 | 2.02 | 1.9 | 17.80 | 18.8 | 3.32 | 3.46 | 42 | 154 | 402 | 2 |
| 155554 | 18 31 50 | 43 56 10 | 2.19 | 1.6 | 37.10 | 32.0 | 3.20 | 3.28 | 68 | 240 | 378 | 2 |
| 155555 | 18 26 50 | 43 56 0 | 2.84 | 1.8 | 21.80 | 21.7 | 2.95 | 3.14 | 56 | 190 | 691 | 3 |
| 155556 | 18 29 50 | 43 55 45 | 4.68 | 3.2 | 17.30 | 16.0 | 3.98 | 4.10 | 64 | 87 | 181 | 126 |

Table 2.--Trace-element concentrations in granite samples from the southeastern Arabian Shield--Continued

| Sample | Y (ppm) | Zr (ppm) | Th/U | eTh/U | K/U | eK/U | Th/K | eTh/eK | K/Rb | Rb/Sr | Al/(Na+K) | Al/(Na+K+Ca) |
|------------------------------|------------|-------------|-------|-------|------|------|-------|--------|--------|--------|-----------|--------------|
| Jabal Mahail (H) | | | | | | | | | | | | |
| 155532 | 32 | 20 | 1.56 | 2.15 | .62 | .67 | 2.50 | 3.21 | 97.70 | 77.75 | 1.33 | 1.27 |
| 155533 | 29 | 22 | 2.94 | 4.22 | 1.13 | 1.26 | 2.60 | 3.34 | 85.73 | 48.14 | 1.36 | 1.29 |
| Jabal Kebad (I) | | | | | | | | | | | | |
| 155534 | 9 | 26 | 1.59 | 1.80 | .63 | .67 | 2.53 | 2.70 | 96.15 | 55.83 | 1.32 | 1.23 |
| 155535 | 1 | 10 | 1.89 | 3.33 | 1.06 | 1.16 | 1.79 | 2.87 | 52.31 | 61.50 | 1.39 | 1.32 |
| Huadbar (J) | | | | | | | | | | | | |
| 155536 | 44 | 133 | 2.06 | 4.09 | .44 | .44 | 4.73 | 9.29 | 120.54 | 35.89 | 1.10 | 1.07 |
| 155537 | 140 | 172 | 2.74 | 3.94 | .11 | .11 | 24.80 | 36.43 | 102.13 | 58.20 | 1.08 | 1.05 |
| 155538 | 64 | 130 | 2.47 | 3.99 | .52 | .53 | 4.74 | 7.58 | 177.68 | 22.80 | 1.09 | 1.03 |
| 155539 | 19 | 70 | 2.46 | 4.19 | .54 | .54 | 4.57 | 7.83 | 126.72 | 4.65 | 1.16 | 1.06 |
| Najran (K) | | | | | | | | | | | | |
| 155540 | 82 | 794 | 3.00 | 2.71 | 1.12 | 1.13 | 2.67 | 2.39 | 365.40 | 1.81 | 1.06 | .93 |
| 155541 | 62 | 663 | 3.16 | 2.29 | .92 | .94 | 3.45 | 2.45 | 367.84 | 1.92 | 1.06 | .94 |
| 155542 | 33 | 115 | 3.60 | 3.38 | .74 | .77 | 4.86 | 4.40 | 339.54 | 1.15 | 1.18 | 1.00 |
| 155543 | 59 | 977 | 7.08 | 5.43 | 3.65 | 3.74 | 1.94 | 1.45 | 603.46 | .67 | 1.09 | .92 |
| Jahal Ashirah (L) | | | | | | | | | | | | |
| 155544 | 22 | 227 | 4.17 | 4.20 | 1.20 | 1.24 | 3.46 | 3.40 | 231.02 | 1.90 | 1.16 | 1.00 |
| 155545 | 26 | 256 | 4.30 | 3.21 | .91 | .91 | 4.84 | 3.52 | 223.82 | 1.71 | 1.17 | 1.01 |
| Jabal al Gaharra (M) | | | | | | | | | | | | |
| 155546 | 66 | 34 | 2.05 | 1.52 | .39 | .40 | 5.24 | 3.77 | 57.67 | 101.00 | 1.14 | 1.09 |
| 155547 | 63 | 49 | 1.35 | 1.15 | .20 | .20 | 6.71 | 5.71 | 48.97 | 217.00 | 1.16 | 1.05 |
| 155548 | 62 | 52 | 2.03 | 1.60 | .26 | .26 | 7.88 | 6.11 | 45.54 | 170.00 | 1.18 | 1.12 |
| 155549 | 35 | 37 | 7.22 | 5.13 | 1.34 | 1.36 | 5.41 | 3.77 | 36.89 | 238.00 | 1.23 | 1.17 |
| Northeast of As Subaykah (N) | | | | | | | | | | | | |
| 155550 | 11 | 55 | <.69 | .67 | .75 | .77 | <.93 | .88 | 409.15 | .65 | 1.24 | 1.12 |
| East of As Subaykah (P) | | | | | | | | | | | | |
| 155551 | 35 | 104 | 1.87 | 1.41 | .69 | .73 | 2.69 | 1.93 | 406.59 | 11.00 | 1.11 | 1.03 |
| 155552 | 25 | 122 | 2.05 | 2.67 | 1.26 | 1.35 | 1.63 | 1.98 | 478.69 | 1.93 | 1.16 | 1.05 |
| Jabal Bani Bwana (Q) | | | | | | | | | | | | |
| 155553 | 50 | 65 | 8.81 | 9.31 | 1.64 | 1.71 | 5.36 | 5.43 | 82.60 | 201.00 | 1.09 | 1.08 |
| 155554 | 141 | 113 | 16.94 | 14.61 | 1.46 | 1.50 | 11.58 | 9.76 | 84.77 | 189.00 | 1.12 | 1.09 |
| 155555 | 9 | 52 | 7.68 | 7.64 | 1.04 | 1.11 | 7.40 | 6.91 | 42.65 | 230.33 | 1.13 | 1.10 |
| 155556 | 28 | 175 | 3.70 | 3.42 | .85 | .88 | 4.35 | 3.90 | 219.69 | 1.44 | 1.26 | 1.11 |

Table 2.—Trace-element concentrations in granite samples from the southeastern Arabian Shield--
Continued

| Sample | Latitude (north) | Longitude (east) | U (ppm) | Ra_{eU} (ppm) | Th (ppm) | e_{Th} (wt %) | K (wt %) | e_K (wt %) | Cu (ppm) | Zn (ppm) | Rb (ppm) | Sr (ppm) |
|---------------------|---------------------|---------------------|--------------|---------------------------|----------------------|---------------------------|---------------|-----------------|---------------|---------------|---------------|---------------|
| Wadi al Habbah (R) | | | | | | | | | | | | |
| 155557 | 18° 30' 40" | 43° 17' 45" | .29 | .3 | <1.40 | .4 | .75 | .78 | 76 | 55 | 19 | 988 |
| 155558 | 18° 31' 45" | 43° 17' 45" | .81 | .7 | 2.80 | 2.4 | .86 | .92 | 79 | 81 | 22 | 803 |
| 155559 | 18° 37' 30" | 43° 17' 15" | .27 | .2 | <1.50 | .4 | 3.04 | 3.01 | 33 | 30 | 43 | 295 |
| Madha (S) | | | | | | | | | | | | |
| 155560 | 18° 39' 55" | 43° 14' 45" | 5.38 | 1.7 | <3.30 | .4 | 1.72 | 1.72 | 26 | 51 | 31 | 633 |
| 155561 | 18° 41' 45" | 43° 12' 50" | .34 | .3 | <1.40 | .3 | 3.69 | 3.50 | 80 | 45 | 52 | 401 |
| Jabal Zayd (T) | | | | | | | | | | | | |
| 155562 | 18° 49' 15" | 43° 19' 45" | 1.52 | 1.4 | 2.80 | 2.3 | 2.61 | 2.66 | 76 | 101 | 49 | 194 |
| 155563 | 18° 49' 30" | 43° 19' 20" | .75 | .7 | 2.10 | 1.9 | 1.98 | 2.03 | 75 | 95 | 37 | 562 |
| Wadi Tafshaf (U) | | | | | | | | | | | | |
| 155564 | 18° 57' 45" | 43° 9' 30" | 1.14 | 1.0 | <1.90 | 1.5 | 3.54 | 3.50 | 70 | 67 | 47 | 13 |
| 155565 | 18° 58' 25" | 43° 10' 15" | 1.36 | 1.3 | 2.60 | 2.0 | 3.36 | 3.34 | 101 | 71 | 36 | 9 |
| 155566 | 18° 59' 55" | 43° 10' 5 | .44 | .5 | <1.60 | .5 | 3.49 | 3.46 | 99 | 62 | 28 | 26 |
| Jabal al Hassir (V) | | | | | | | | | | | | |
| 155567 | 19° 40' 45" | 43° 6' 40" | 3.31 | 3.3 | 8.90 | 8.4 | 4.17 | 4.12 | 101 | 141 | 82 | 15 |
| 155568 | 19° 39' 0 | 43° 5' 40" | 3.52 | 3.0 | 9.73 | 7.8 | 4.25 | 4.41 | 102 | 115 | 69 | 18 |
| 155569 | 19° 36' 0 | 43° 5' 20" | 6.36 | 6.2 | 14.90 | 13.9 | 3.96 | 3.96 | 81 | 197 | 154 | 29 |
| 155570 | 19° 30' 0 | 43° 3' 55" | 6.55 | 5.8 | 18.00 | 17.2 | 3.84 | 3.92 | 76 | 75 | 174 | 60 |
| 155571 | 19° 25' 25" | 43° 4' 35" | 2.92 | 2.8 | 9.75 | 8.9 | 4.29 | 4.45 | 75 | 155 | 63 | 31 |
| 155572 | 19° 23' 20" | 43° 4' 10" | 5.24 | 4.2 | 10.10 | 9.8 | 4.56 | 4.68 | 111 | 91 | 112 | 58 |

Table 2.—Trace-element concentrations in granite samples from the southeastern Arabian Shield—
Continued

| Sample | γ (ppm) | Zr (ppm) | Th/U | eTh/U | K/U | eK/U | Th/K | eTh/ek | ek/Rb | Rb/Sr | $Al/(Na+K)$ |
|---------------------|-------------------|---------------|--------|---------|-------|--------|--------|----------|----------|---------|-------------|
| Wadi al Habbah (R) | | | | | | | | | | | |
| 155557 | <1 | 7.6 | <4.83 | 1.38 | 2.58 | 7.69 | <1.87 | .51 | 393.23 | .02 | 1.83 |
| 155558 | 6 | 5.8 | 3.46 | 2.96 | 1.07 | 1.14 | 3.24 | 2.61 | 392.44 | .03 | 1.97 |
| 155559 | <1 | 2.2 | <5.56 | 1.48 | 11.25 | 11.15 | <4.9 | .13 | 706.60 | .15 | 1.35 |
| Madha (S) | | | | | | | | | | | |
| 155560 | <1 | 5.4 | <6.61 | .07 | .32 | 10.20 | <1.92 | .23 | 554.33 | .05 | 1.56 |
| 155561 | 12 | 3.2 | <4.12 | .88 | 10.84 | <.38 | .09 | 708.83 | .13 | 1.34 | 1.12 |
| Jabal Zayd (T) | | | | | | | | | | | |
| 155562 | 2 | 32 | 1.84 | 1.51 | 1.71 | 1.75 | 1.07 | .86 | 531.98 | .25 | 1.40 |
| 155563 | 3 | 5.6 | 2.80 | 2.53 | 2.65 | 2.71 | 1.06 | .94 | 536.24 | .07 | 1.61 |
| Wadi Tafshaf (U) | | | | | | | | | | | |
| 155564 | 33 | 61 | <1.67 | 1.32 | 3.10 | 3.07 | <.54 | .43 | 752.44 | 3.62 | 1.21 |
| 155565 | 38 | 45 | 1.91 | 1.47 | 2.47 | 2.46 | .77 | .60 | 935.93 | 4.00 | 1.20 |
| 155566 | 28 | 33 | <3.64 | 1.14 | 7.92 | 7.86 | <.46 | .14 | 1,245.24 | 1.08 | 1.20 |
| Jabal al Hassir (V) | | | | | | | | | | | |
| 155567 | 37 | 490 | 2.69 | 2.54 | 1.26 | 1.24 | 2.14 | 2.04 | 508.22 | 5.47 | 1.05 |
| 155568 | 29 | 531 | 2.76 | 2.22 | 1.21 | 1.25 | 2.29 | 1.77 | 616.00 | 3.83 | .99 |
| 155569 | 77 | 347 | 2.34 | 2.19 | .62 | .62 | 3.76 | 3.51 | 257.13 | 5.31 | 1.02 |
| 155570 | 17 | 98 | 2.75 | 2.63 | .50 | .60 | 4.68 | 4.39 | 220.90 | 2.90 | 1.11 |
| 155571 | 30 | 600 | 3.34 | 3.05 | 1.47 | 1.52 | 2.27 | 2.00 | 681.26 | 2.03 | 1.23 |
| 155572 | 29 | 147 | 1.93 | 1.87 | .87 | .89 | 2.22 | 2.09 | 406.93 | 1.93 | 1.10 |

disequilibrium between ^{238}U and ^{226}Ra in granitic rocks is common (for example, Stuckless and others, 1977), and, therefore, U is used in preference to RaeU for radioelement ratios and statistical calculations. Accuracy and precision (2σ) for RaeU values are within \pm the quantity (2 percent of the amount reported plus 0.1 ppm absolute).

Concentrations of copper (Cu), zinc (Zn), rubidium (Rb), strontium (Sr), yttrium (Y), and zirconium (Zr) (table 2) were determined by X-ray fluorescence on loose, finely ground (-200 mesh) sample powders. Estimated precision is \pm 10 percent (2σ) for Rb, Sr, Y, and Zr, and \pm 20 percent for Cu and Zn; however, Cu values may not be accurate within the limits of precision because standard values at low concentrations are not well known. Niobium and molybdenum were looked for, but concentrations were generally below the limits of detection (6 and 4 ppm, respectively), and, therefore, results for these elements are not reported.

Statistical formulae and methods used are those described by VanTrump and Miesch (1977). Statistical parameters: mean, standard deviation, skewness, kurtosis, and correlation, for trace elements (those reported in parts per million in table 2) were calculated from the logarithms of the data. Antilog values are reported for geometric means and standard deviations. Normative mineralogy (table 1) was calculated by methods described by Stuckless and VanTrump (1979).

RESULTS AND DISCUSSION

Petrogenetic considerations

The granitic samples from the southeastern Arabian Shield yield a large range of chemical compositions as shown by minimum and maximum values and by large standard deviations (table 3). Although the mean and range of major- and trace-element contents are similar to those reported for postorogenic granites from the northeastern part of the Shield (Stuckless and others, *in press*; 1981b) standard deviations for trace-element contents and ratios are much larger. Also, the fact that absolute values for skewness and kurtosis are larger for the trace-element contents of granites from the southeastern part of the Shield indicates a poorer fit to a log-normal distribution.

The granites sampled in the current study can be subdivided into three groups on the basis of field occurrence. The granites from Wadi al Habbah and Madha (R and S, respectively, fig. 1) are largely concordant bodies with regional metamorphic structures and, therefore, may be orogenic in origin. The granites of Najran, Jabal Ashirah, and Jabal al

Table 3.--Statistical summary for chemical compositions of granite samples from the southeastern Arabian Shield
 [Oxide concentrations are in weight percent; trace-element contents are in parts per million. Average granite values from Krauskopf (1967) with major elements converted to oxides and normalized to 100 percent]

| | Number of samples | Minimum | Maximum | Mean | Standard deviation | Skewness | Kurtosis | Average granite |
|--------------------------------|----------------------|---------|---------|-------|-----------------------|----------|----------|--------------------|
| SiO ₂ | 73 | 66.66 | 80.19 | 74.81 | +2.35 | -1.25 | 2.32 | 72.64 |
| Al ₂ O ₃ | 73 | 10.89 | 17.24 | 13.41 | +1.13 | 0.73 | 1.28 | 13.67 |
| Fe ₂ O ₃ | 73 | 0.36 | 5.29 | 1.38 | +0.91 | 2.00 | 5.12 | 3.63 |
| MgO | 73 | 0.05 | 1.14 | 0.15 | +0.18 | 3.30 | 13.32 | 0.25 |
| CaO | 73 | 0.12 | 4.58 | 0.76 | +0.73 | 3.22 | 12.64 | 2.10 |
| Na ₂ O | 73 | 3.05 | 5.14 | 3.99 | +0.46 | 0.23 | -0.30 | 3.54 |
| K ₂ O | 73 | 0.90 | 6.48 | 4.28 | +0.88 | -1.49 | 4.81 | 3.74 |
| TiO ₂ | 73 | 0.01 | 0.48 | 0.09 | +0.10 | 1.62 | 2.67 | 0.36 |
| P ₂ O ₅ | 73 | 0.01 | 0.13 | 0.03 | +0.03 | 2.19 | 3.80 | 0.03 |
| MnO | 73 | 0.01 | 0.17 | 0.04 | +0.03 | 1.71 | 3.88 | 0.05 |
| U | 73 | 0.27 | 26.9 | 4.46 | +7.61 -2.81 | -0.88 | 0.93 | 4.8 |
| RaeU | 73 | 0.2 | 31.0 | 4.16 | +7.32 -2.65 | -0.86 | 0.85 | - |
| Th | 66 | 2.1 | 73.7 | 16.7 | +18.5 -8.8 | -0.72 | 0.74 | 17 |
| eTh | 73 | 0.3 | 106 | 12.2 | +30.1 -8.7 | -1.49 | 2.07 | - |
| Cu | 73 | 22 | 148 | 59.9 | +30.3 -20.1 | -0.49 | -0.11 | 10 |
| Zn | 73 | 30 | 240 | 92.6 | +49.1 -32.1 | -0.04 | -0.14 | 40 |
| Rb | 73 | 19.0 | 952 | 195 | +332 -123 | -0.45 | -0.67 | 150 |
| Sr | 72 | 1.0 | 988 | 24.0 | +108.6 -19.7 | 0.25 | -0.80 | 285 |
| Y | 70 | 1.0 | 221 | 38.2 | +77.7 -25.6 | -0.92 | 1.03 | 40 |
| Zr | 73 | 10.0 | 977 | 95.4 | +128.5 -54.8 | 0.43 | 1.07 | 180 |
| K/Rb | 73 | 36.9 | 1250 | 176 | +265 -106 | 0.09 | -1.11 | - |
| Rb/Sr | 72 | 0.02 | 372 | 8.11 | +92.9 -7.46 | -0.35 | -0.68 | - |
| eTh/U | 73 | 0.10 | 14.6 | 2.74 | +2.68 -1.36 | -2.01 | 10.17 | - |
| eTh/eK | 73 | 0.10 | 36.4 | 3.48 | +7.22 -2.35 | -1.35 | 2.28 | - |
| eK/U | 73 | 0.11 | 11.2 | 0.79 | +1.16 -0.47 | 0.62 | 1.00 | - |

Hassir (K, L, and V, respectively, fig. 1) are similar to postorogenic peralkaline intrusive rocks in the northeastern part of the Shield in that rim phases are very coarse grained, one-feldspar granites in which black amphibole is the dominant mafic mineral. However, these three granites from the southeastern part of the Shield have molar $\text{Al}/(\text{Na}+\text{K})$ ratios greater than 1 (table 2) and are not peralkaline by the definition of Shand (1951) but rather are evenly divided between metaluminous and weakly peraluminous. These plutons are referred to as postorogenic hypersolvus granites. The remaining plutons are referred to as postorogenic subsolvus granites; all samples from these plutons have molar ratios $\text{Al}/(\text{Na}+\text{K}+\text{Ca})$ greater than 1 and are, therefore, peraluminous (table 2).

Mean values and standard deviations for oxide and trace-element contents and ratios of samples from the southeastern Shield collected for this study are given in table 4 for the three groups of granites and for all of the postorogenic granites combined. The data clearly demonstrate that the concordant and possibly orogenic granites are chemically distinct from the postorogenic granites in that they contain more aluminum, magnesium, calcium, and strontium and less silica, potassium, uranium, thorium, rubidium, and yttrium. All of the trace-element ratios (table 4) and the Na/K ratio are markedly different for the two groups, and it is unlikely that they are related genetically.

Differences also exist between the hypersolvus and subsolvus postorogenic granites, but these are generally smaller than the differences noted between the postorogenic and possibly orogenic granites. The average iron content of the hypersolvus granites is more than double that of the subsolvus granites (table 4). The most striking difference is in zirconium content, which is more than four times greater in the hypersolvus granites. This high mean value (343 ppm, table 4) is similar to that noted for metaluminous postorogenic granites from the northeastern part of the Shield (381 ppm, Stuckless and others, 1982b), and is distinctly anomalous relative to the approximately 100 ppm saturation limits for zirconium cited by Watson (1979) for experimental systems that have similar molar ratios of $\text{Al}/(\text{Na}+\text{K})$. Both iron and zirconium enrichments are commonly associated with silicic peralkaline igneous rocks. The hypersolvus granites also differ from the other postorogenic granites from the southeastern part of the Shield in trace-element ratios. The mean K/Rb ratio is larger and the mean Rb/Sr ratio is smaller in the hypersolvus granites (table 4). These differences are similar to those noted between metaluminous and peralkaline or peraluminous samples from single plutons within the northeastern part of the Shield (Stuckless and others, 1982b), and can therefore be attributed to differing degrees of magma

Table 4.--Comparison of means and standard deviations for subgroupings of granite samples from the southeastern Arabian Shield

| | Postorogenic samples | | | | | | | |
|--------------------------------|----------------------|--------------------|---------------------|--------------------|-------------------|--------------------|---------------------------|--------------------|
| | All samples | | Hypersolvus samples | | Subsolvus samples | | Possible orogenic samples | |
| | Mean | Standard deviation | Mean | Standard deviation | Mean | Standard deviation | Mean | Standard deviation |
| SiO ₂ | 75.02 | ±2.20 | 72.73 | ±2.91 | 75.52 | ±1.68 | 71.84 | ±2.56 |
| Al ₂ O ₃ | 13.24 | ±0.93 | 13.20 | ±0.93 | 13.24 | ±0.94 | 15.81 | ±0.91 |
| Fe ₂ O ₃ | 1.40 | ±0.92 | 2.69 | ±1.31 | 1.12 | ±0.48 | 1.18 | ±0.87 |
| MgO | 0.13 | ±0.13 | 0.13 | ±0.09 | 0.13 | ±0.14 | 0.42 | ±0.46 |
| CaO | 0.62 | ±0.41 | 0.80 | ±0.31 | 0.58 | ±0.04 | 2.62 | ±1.45 |
| Na ₂ O | 3.96 | ±0.45 | 3.89 | ±0.35 | 3.97 | ±0.47 | 4.48 | ±0.32 |
| K ₂ O | 4.42 | ±0.64 | 5.06 | ±0.46 | 4.28 | ±0.59 | 2.42 | ±1.58 |
| TiO ₂ | 0.09 | ±0.10 | 0.21 | ±0.13 | 0.06 | ±0.07 | 0.11 | ±0.12 |
| P ₂ O ₅ | 0.03 | ±0.03 | 0.03 | ±0.02 | 0.03 | ±0.03 | 0.05 | ±0.05 |
| MnO | 0.04 | ±0.03 | 0.05 | ±0.02 | 0.04 | ±0.04 | 0.04 | ±0.05 |
| U | 5.14 | +6.53 -2.88 | 3.93 | +2.13 -1.38 | 5.44 | +7.60 -3.17 | 0.65 | +1.64 -0.47 |
| RaeU | 4.89 | +6.24 -2.74 | 3.71 | +1.69 -1.16 | 5.19 | +7.33 -3.04 | 0.46 | +0.63 -0.27 |
| Th | 17.13 | +18.02 -8.78 | 12.74 | +4.18 -3.15 | 18.31 | +21.27 -9.84 | -- | -- |
| eTh | 15.39 | +22.98 -9.22 | 11.17 | +3.83 -2.85 | 16.48 | +27.69 -10.33 | 0.54 | +0.71 -0.31 |
| Cu | 60.4 | +29.9 -20.0 | 72.2 | +44.9 -27.9 | 58.1 | +26.6 -18.2 | 52.8 | +38.3 -22.2 |
| Zn | 97.0 | +47.0 -31.7 | 114 | +46 -33 | 93.7 | +46.1 -30.9 | 49.8 | +21.5 -15.0 |
| Rb | 224 | +317 -131 | 113 | +51 -35 | 259 | +373 -153 | 31 | +17 -11 |
| Sr | 18.9 | +67.8 -14.8 | 51.8 | +51.1 -25.7 | 15.2 | +57.7 -12.0 | 569 | +366 -223 |
| Y | 39.9 | +79.2 -26.5 | 37.2 | +24.5 -14.8 | 40.5 | +92.0 -28.1 | 8.5 | +5.4 -3.3 |
| Zr | 101.0 | +134.9 -57.8 | 343.3 | +401.7 -185.1 | 77.7 | +63.9 -35.1 | 44.1 | +28.6 -17.4 |
| K/Rb | 162 | +235 -96 | 372 | +189 -125 | 136 | +189 -79 | 533 | +183 -136 |
| Rb/Sr | 11.8 | +92.6 -10.5 | 2.2 | +1.8 -1.0 | 17.0 | +141 -15.2 | 0.05 | +0.08 -0.03 |
| eTh/U | 2.99 | +2.02 -1.21 | 2.85 | +1.00 -0.70 | 3.03 | +2.23 -1.29 | 0.83 | +2.60 -0.63 |
| eTh/eK | 4.15 | +5.96 -2.45 | 2.62 | +1.16 -0.80 | 4.58 | +7.14 -2.79 | 0.32 | +0.92 -0.24 |
| eK/U | 0.72 | +0.88 -0.40 | 1.09 | +0.68 -0.42 | 0.66 | +0.85 -0.37 | 2.57 | +9.04 -0.63 |

evolution such that the subsolvus granites represent more evolved members of a single petrogenetic suite, all of which is highly evolved.

Permissive evidence for a genetic relationship among all the postorogenic granites is provided by plots of the data in the normative feldspar system (fig. 2) and the normative Q-Ab-Or (granite) system (fig. 3). The fact that most of the data for both the hypersolvus and subsolvus granites (fig. 2) cluster along the 2-kb water-saturated boundary curve suggests equilibration under similar physical conditions. Significantly, data for three of the five possibly orogenic granite samples plot far from equilibrium feldspar compositions within the area where plagioclase is expected to crystallize first.

The trend of data within the Q-Ab-Or system (fig. 3) is similar to that noted for single postorogenic plutons in the northeastern part of the Shield (Stuckless, unpublished data). Data for several samples of hypersolvus granite plot well below the water-saturated, polybaric minimum and form a trend directly toward the quartz apex. This trend can be interpreted to reflect a high-pressure partial melting at water-undersaturated conditions and crystal fractionation at lower pressure and increasing water activity. As in the feldspar system, more than half of the data for the possibly orogenic granites plot well away from the data for the postorogenic granites.

Correlation analysis (table 5) also yields permissive evidence for a genetic relationship between the two types of postorogenic granites. Several elemental pairs that typically covary in single petrogenetic suites exhibit high correlation coefficients (for example, CaO or MgO with Sr and TiO₂ with Fe₂O₃) in spite of the wide geographical separation of plutons. This fact suggests that the various plutons were derived from similar source material by similar processes and that differentiation followed similar paths. However, as would be predicted from geographical consideration, several low correlation coefficients (such as K₂O with most variables) preclude a true cogenetic relationship.

Economic considerations

Concentration data indicate at least a moderate favorability for mineralized areas in the southeastern Arabian Shield. Copper, zinc, uranium, and yttrium each occur in at least one pluton in concentrations that are well above average levels for granite. Although individual values for copper (and to a lesser extent zinc) have large analytical uncertainties, the average concentrations of these elements

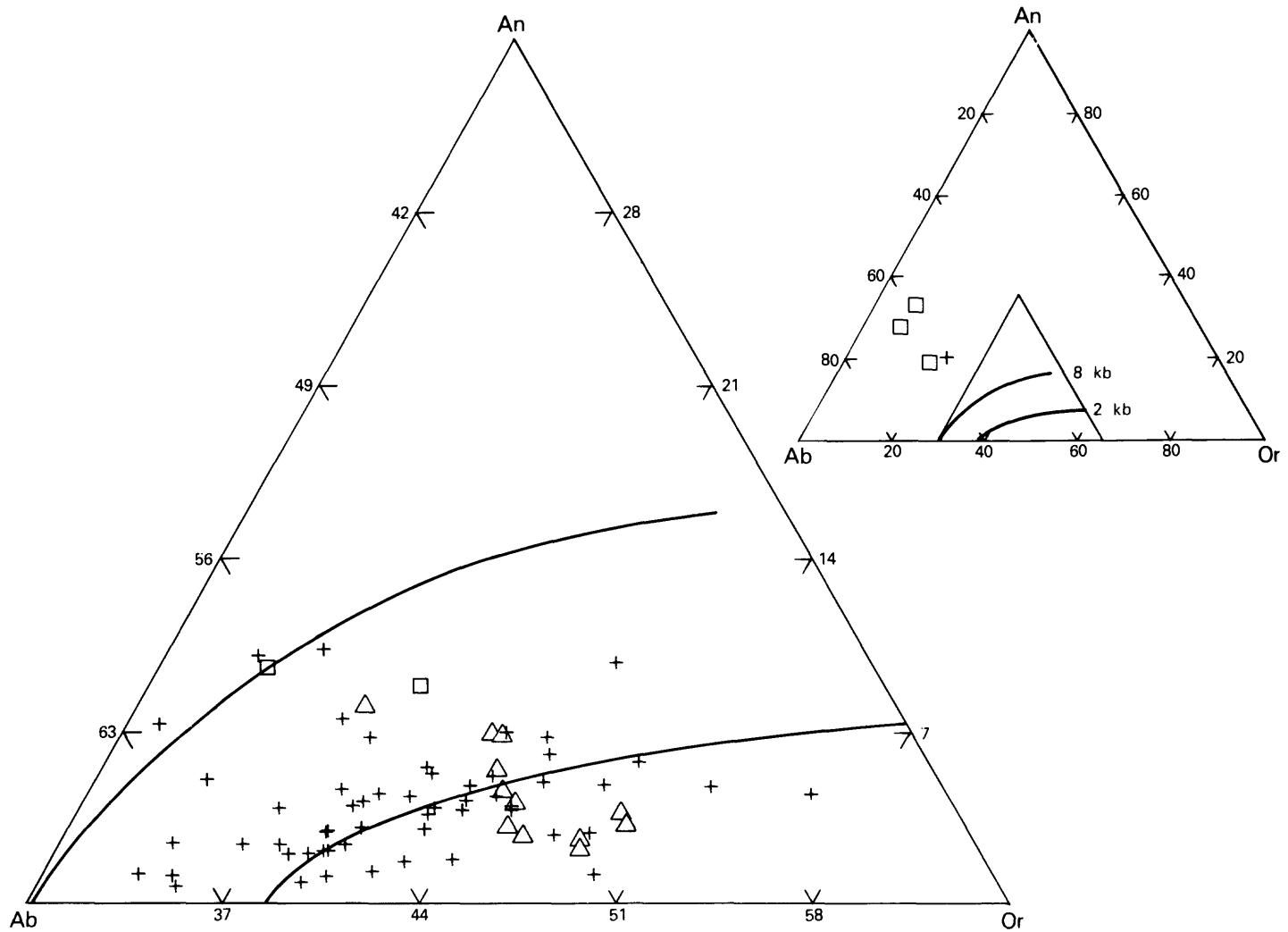


Figure 2.--Ternary diagram showing normative feldspar data for granite samples from the southeastern Arabian Shield; An, anorthite, Ab, albite, Or, orthoclase. Curved lines show the position of water-saturated, eutectic compositions at 2 and 8 kb (Whitney, 1975). Squares represent possible orogenic granites, triangles represent postorogenic hypersolvus granites, and pluses represent postorogenic subsolvus granites.

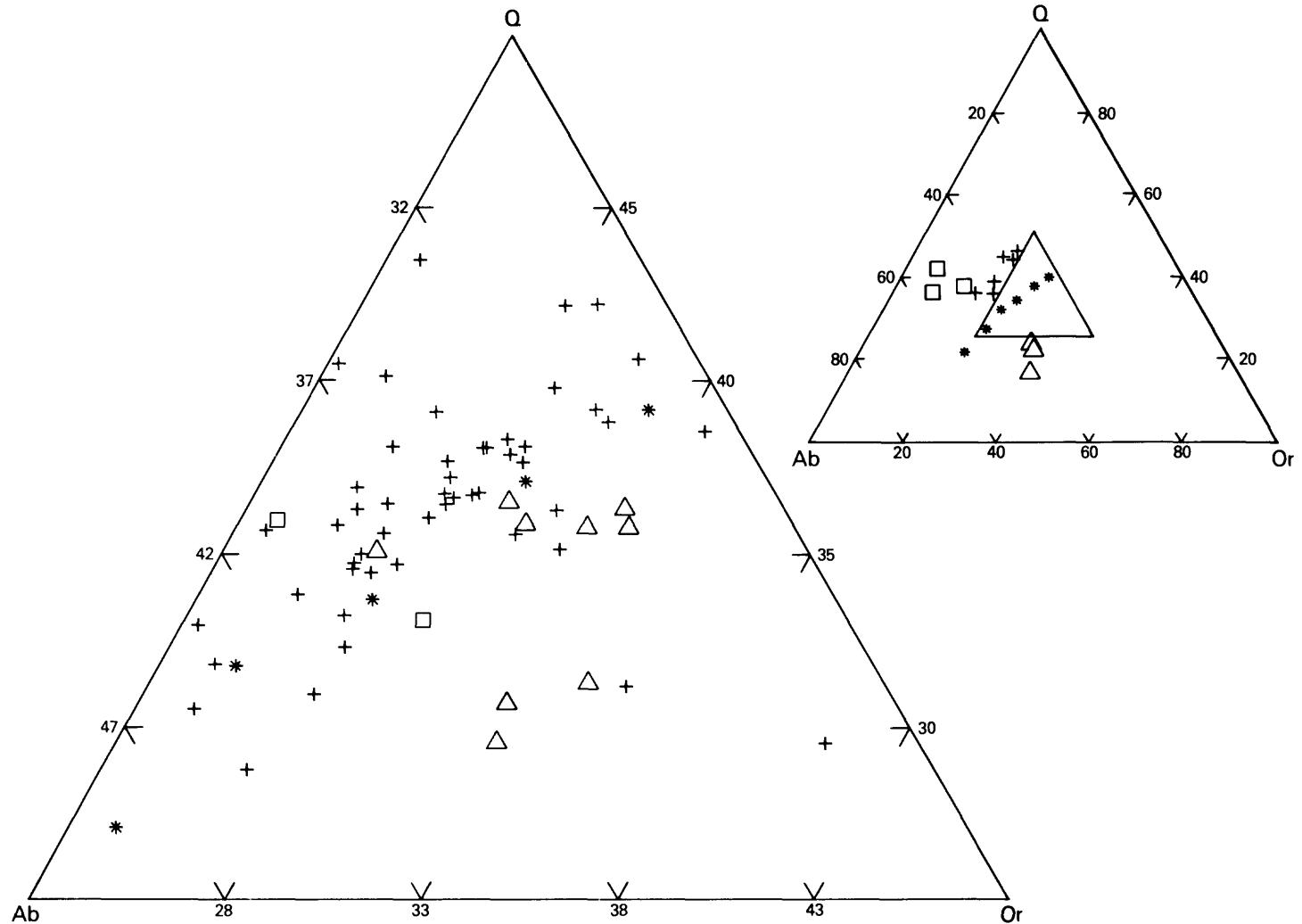


Figure 3.--Ternary diagram showing normative quartz (Q), albite (Ab), and orthoclase (Or) data for granite samples from the southeastern Arabian Shield. The position of the polybaric ternary minimum (Tuttle and Bowen, 1958; Luth and others, 1964) is shown by asterisks. Squares represent possible orogenic granites, triangles represent postorogenic hypersolvus granites, and pluses represent postorogenic subsolvus granites.

Table 5.--Correlation matrix for postorogenic granite samples from the southeastern Arabian Shield and correlation of selected variables with latitude and longitude
 [Oxides are treated arithmetically. Elements and ratios are treated logarithmically. Numbers below the diagonal blank space indicate the number of samples used in the correlation analysis.
 Abbreviations are defined in text]

| | SiO ₂ | Al ₂ O ₃ | Fe ₂ O ₃ | MgO | CaO | K ₂ O | TiO ₂ | U | Y | eTh | Rb | Sr | Zr | Rb/Sr | K/Rb | eTh/U | eTh/eK | eK/U |
|--------------------------------|------------------|--------------------------------|--------------------------------|-------|-------|------------------|------------------|-------|-------|-------|-------|-------|-------|-------|-------|-------|--------|-------|
| SiO ₂ | -0.71 | -0.68 | -0.55 | 0.73 | -0.29 | -0.30 | 0.79 | 0.32 | 0.25 | 0.34 | 0.34 | -0.62 | -0.43 | 0.57 | -0.37 | 0.10 | 0.39 | -0.37 |
| Al ₂ O ₃ | 68 | 0.05 | 0.49 | 0.53 | 0.58 | -0.20 | 0.24 | -0.40 | -0.48 | -0.48 | -0.15 | 0.41 | -0.27 | -0.35 | 0.10 | -0.21 | -0.45 | 0.37 |
| Fe ₂ O ₃ | 68 | 68 | 0.27 | 0.47 | -0.13 | 0.47 | 0.88 | -0.18 | 0.01 | -0.11 | -0.36 | 0.43 | 0.73 | -0.44 | 0.43 | 0.08 | -0.19 | 0.26 |
| MgO | 68 | 68 | 68 | 0.88 | 0.0 | 0.15 | 0.55 | -0.33 | -0.48 | -0.26 | -0.40 | 0.76 | 0.13 | -0.69 | 0.36 | 0.06 | -0.22 | 0.29 |
| CaO | 68 | 68 | 68 | 68 | 0.03 | -0.05 | 0.68 | -0.31 | -0.36 | -0.33 | -0.50 | 0.75 | 0.28 | -0.73 | 0.47 | -0.09 | -0.32 | 0.29 |
| Na ₂ O | 68 | 68 | 68 | 68 | 68 | -0.47 | -0.10 | -0.26 | -0.11 | -0.24 | 0.07 | -0.16 | -0.23 | 0.14 | -0.15 | 0.02 | -0.17 | 0.18 |
| K ₂ O | 68 | 68 | 68 | 68 | 68 | 68 | 0.43 | 0.22 | 0.32 | 0.19 | -0.04 | 0.22 | 0.65 | -0.16 | 0.20 | 0.0 | 0.04 | -0.04 |
| TiO ₂ | 68 | 68 | 68 | 68 | 68 | 68 | 68 | -0.29 | -0.17 | 0.19 | -0.47 | 0.66 | 0.75 | -0.65 | 0.53 | 0.12 | -0.25 | 0.36 |
| U | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 0.61 | 0.83 | 0.69 | -0.37 | 0.10 | 0.54 | -0.64 | -0.12 | 0.81 | -0.98 |
| Y | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 0.59 | 0.46 | -0.54 | 0.31 | 0.57 | -0.39 | 0.07 | 0.54 | -0.57 |
| eTh | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 0.76 | -0.39 | 0.16 | 0.59 | -0.71 | 0.45 | 0.99 | -0.81 | |
| Rb | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | -0.61 | -0.31 | 0.84 | -0.99 | 0.26 | 0.78 | -0.71 | |
| Sr | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 0.31 | -0.95 | 0.63 | -0.11 | -0.43 | 0.40 |
| Zr | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | -0.34 | 0.41 | 0.21 | 0.17 | -0.06 | |
| Rb/Sr | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | 67 | -0.84 | 0.18 | 0.62 | -0.57 | |
| K/Rb | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | -0.25 | -0.76 | 0.68 | | |
| eTh/U | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 0.46 | 0.13 | |
| eTh/eK | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | -0.81 | -0.81 | |
| eK/U | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | 68 | |
| lat. | 0.47 | -0.37 | -0.35 | -0.07 | -0.16 | -0.45 | 0.04 | -0.30 | 0.49 | 0.30 | 0.54 | 0.36 | -0.21 | -0.14 | 0.29 | -0.35 | 0.18 | 0.54 |
| long. | 0.13 | -0.10 | -0.07 | -0.14 | -0.10 | -0.17 | 0.0 | -0.16 | 0.47 | 0.40 | 0.56 | 0.59 | -0.32 | -0.18 | 0.47 | -0.58 | 0.25 | 0.57 |

for all of the postorogenic plutons exceed average concentrations for granitic rock ($\text{Cu}=10$ ppm, $\text{Zn}=40$ ppm, Krauskopf, 1967) by factors of approximately 6 and 2.5, respectively. Furthermore, semiquantitative spectrographic analyses for these elements in more than 600 granitic samples from the southeastern Shield also yield anomalous average values (du Bray and others, 1983).

In addition to generally anomalous levels of copper and zinc in the postorogenic granites in the southeastern Arabian Shield, there are a few plutons for which chalcophile-element contents are distinctly greater than the anomalously high regional background levels. Even though the low number of samples and large analytical uncertainties preclude a statistically valid assessment of average copper and zinc contents for individual plutons, it is noteworthy that average copper contents at Jabal Kebad and Jabal al Hassir (locations I and V, fig. 1) are more than one standard deviation greater than the regional background. The average zinc content at Jabal Bani Bwana (location Q, fig. 1) is, likewise, more than one standard deviation greater than the regional background. Average zinc and copper contents from several other areas, although not anomalous relative to the regional high background, are distinctly high relative to average granite: Jabal Tarban, Jabal Sabhah, Jabal Hawshat ibn Hawayl, Jabal Sahah, Najran, and Jabal al Gaharra (locations A, C, D, F, K, and M, respectively, fig. 1).

Yttrium is geochemically very similar to the rare-earth elements (REE) dysprosium and holmium (Felsche and Herrmann, 1978). Yttrium, therefore, provides an accurate prediction of at least middle REE contents. The average yttrium content (40 ppm, table 4) for the postorogenic granites from the southeastern Arabian Shield is identical to that listed by Krauskopf (1967) for granites in general, and it is, therefore, not considered to be anomalous. However, average yttrium content (156 ppm) for samples from Jabal Sabhah (location C, fig. 1) is greater than the regional mean value plus one standard deviation (119, tables 2 and 4). Samples from Jabal Hawshat ibn Hawayl (location D, fig. 1) yield a mean yttrium content of 111 ppm, which is just slightly less than the regional mean plus one standard deviation. These anomalous values suggest the possibility of REE mineralization in the northern part of the southeastern Shield. In addition to these larger scale anomalies, there are a few anomalous samples. One of these, sample 155537, was collected along the margin of a pegmatite. Such pegmatites could provide sites for significant REE concentrations.

Average uranium and thorium contents for postorogenic granites from the southeastern Shield (table 4) are nearly identical to those reported for approximately 2,500 granitic samples from the contiguous United States ($U = 3.54^{+4.58}_{-2.00}$, $Th = 16.76^{+22.47}_{-9.60}$, Stuckless and VanTrump, 1982) and are slightly lower than those reported for postorogenic granites of the northeastern Arabian Shield ($U = 5.63^{+4.08}_{-2.36}$, $Th = 16.54^{+11.00}_{-6.61}$, Stuckless and others, *in press*). Thus, the postorogenic granites of the southeastern Shield do not indicate a region that has an anomalous endowment of radioelements. However, like yttrium, uranium and to a lesser extent thorium occur in anomalous amounts in a few plutons.

A comparison of uranium content in granite samples from the southeastern Shield with normal contents (the mean plus or minus one standard deviation) of granites from the contiguous United States was plotted as a function of areal distribution (fig. 4). Five plutons (Jabal Tarban, Jabal Sabhah, Jabal Hawshat ibn Hawayl, Jabal al Hawshaw, and Jabal Sahah) from the northern part of the southeastern Shield (locations A, C, D, E, and F) and the pluton at Jabal al Gaharra (location M) exhibit anomalously high uranium contents. At least four of these six plutons, Jabal Tarban, Jabal Sabhah, Jabal Hawshat ibn Hawayl, and Jabal al Gaharra, are also anomalous with respect to other elements of economic interest.

The areal distribution of thorium values relative to the mean and standard deviation of thorium values in granites from the contiguous United States (fig. 5) shows that thorium content is anomalously high at Jabal Sabhah and Jabal al Hawshaw (locations C and E) and that only one other sample is anomalously rich in thorium. This sample (155537, location J) is a distinctly banded rock that was collected next to a pegmatite at Huqban and is, therefore, probably not representative of the pluton as a whole. This sample is also anomalously enriched in uranium and yttrium. These elements could also have been introduced from a late-stage pegmatite that was anomalously rich in incompatible elements.

Stuckless and VanTrump (1982) have pointed out that Th/U ratios can be useful in the assessment of uranium favorability for secondary deposits. Very large or highly variable Th/U ratios may indicate a significant post-crystallization loss of uranium from granite. The mean Th/U ratio for the postorogenic granites (2.99, table 4) is much smaller than that reported for the average granite in the United States (4.73, Stuckless and VanTrump, 1982), and the standard deviation for the postorogenic granites (+2.02, -1.21, table 4) is much smaller than that for the United States granites (+5.97, -2.64). Therefore, the data do not suggest any significant amount of post-magmatic mobility of uranium. This conclusion is supported by the high correlation coefficient for

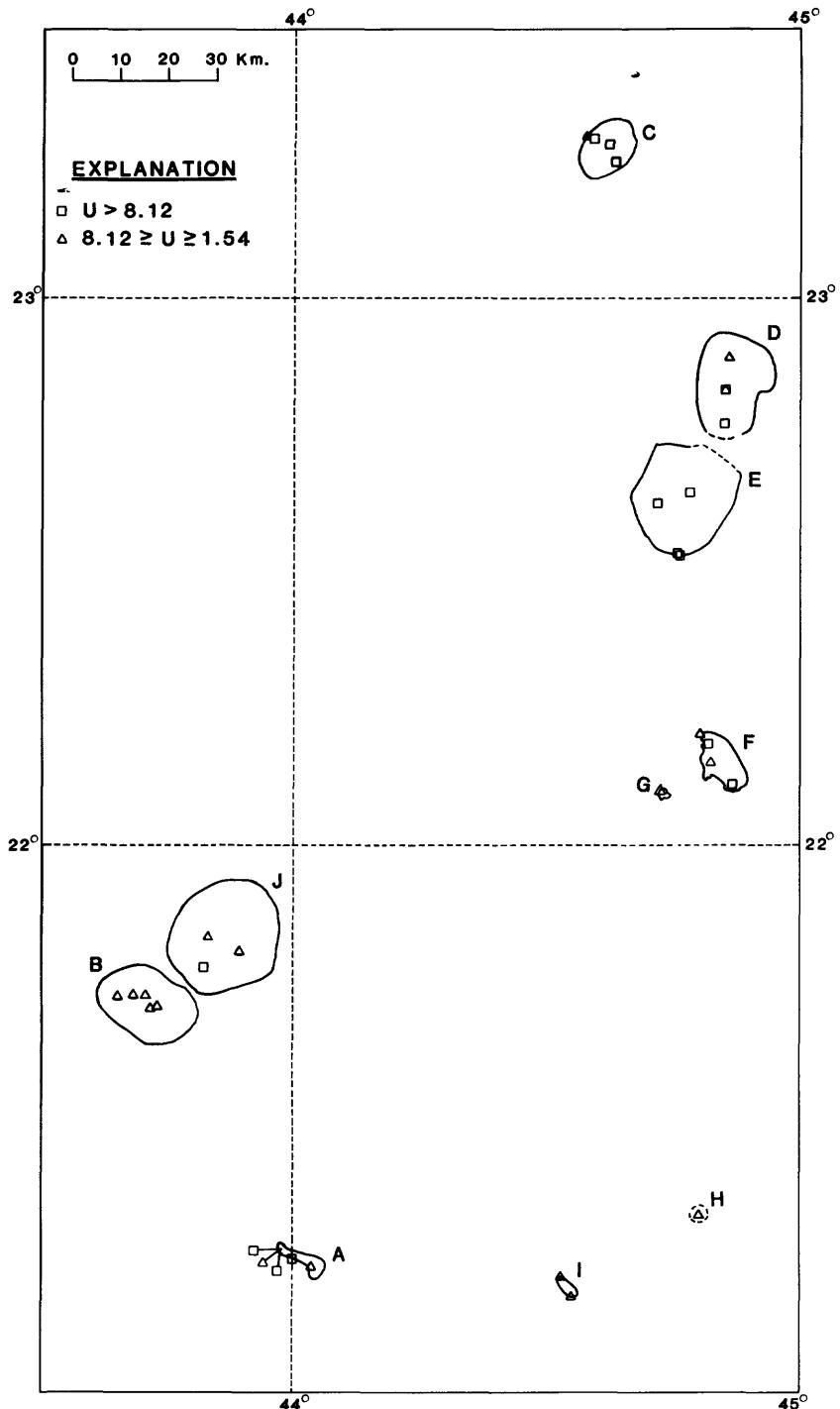


Figure 4.--Map showing areal distribution of uranium values for granite samples from the southeastern part of the Arabian Shield; sample locations and approximate pluton outlines also shown. Symbols represent uranium contents in parts per million relative to the mean and standard deviation of uranium contents in approximately 2,500 granitic samples from the contiguous United States (Stuckless and VanTrump, 1982): triangles, samples having uranium contents within a normal range for granites; squares, anomalously uraniferous samples; circles, anomalously uranium-poor samples.

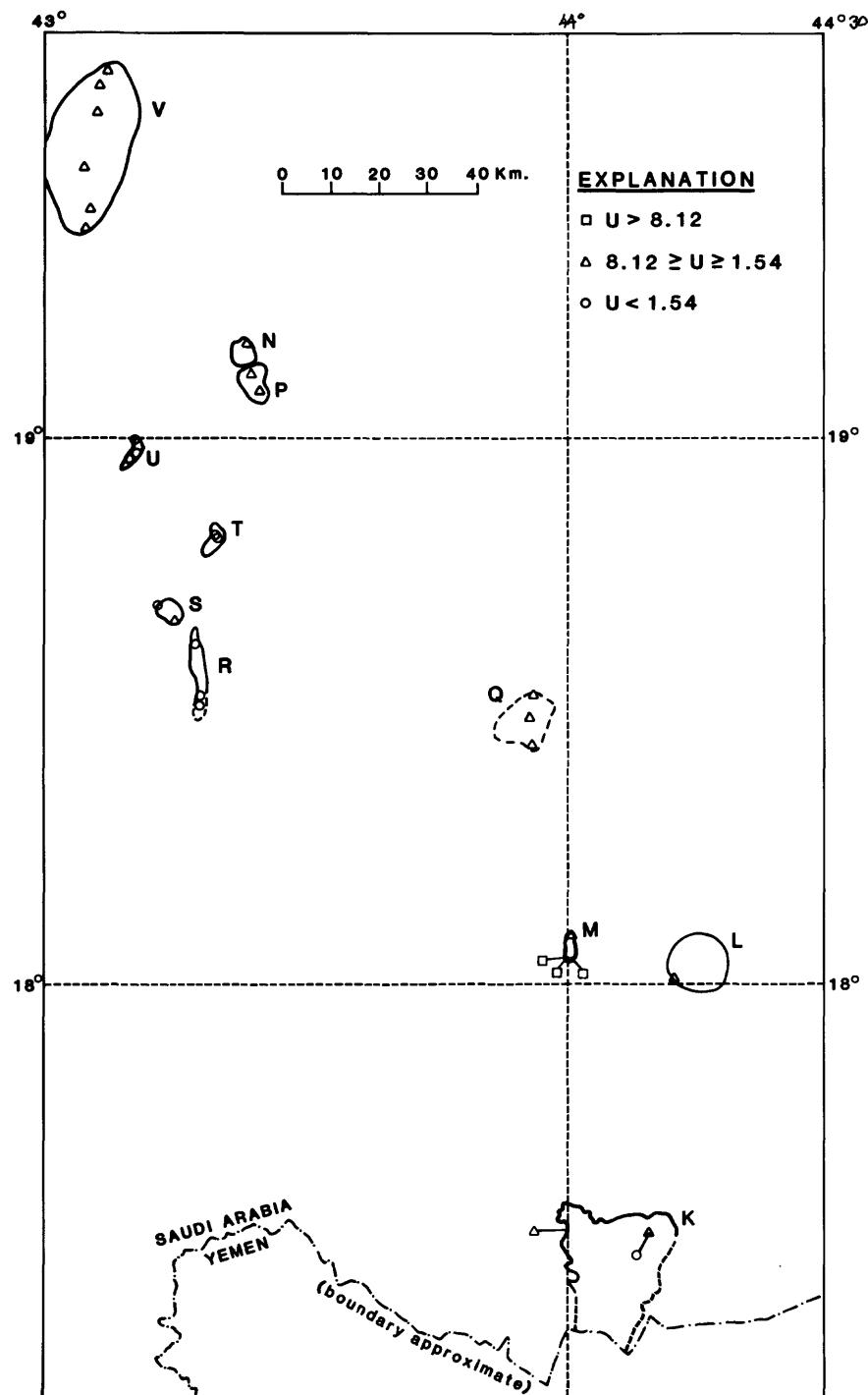


Figure 4.--Continued.

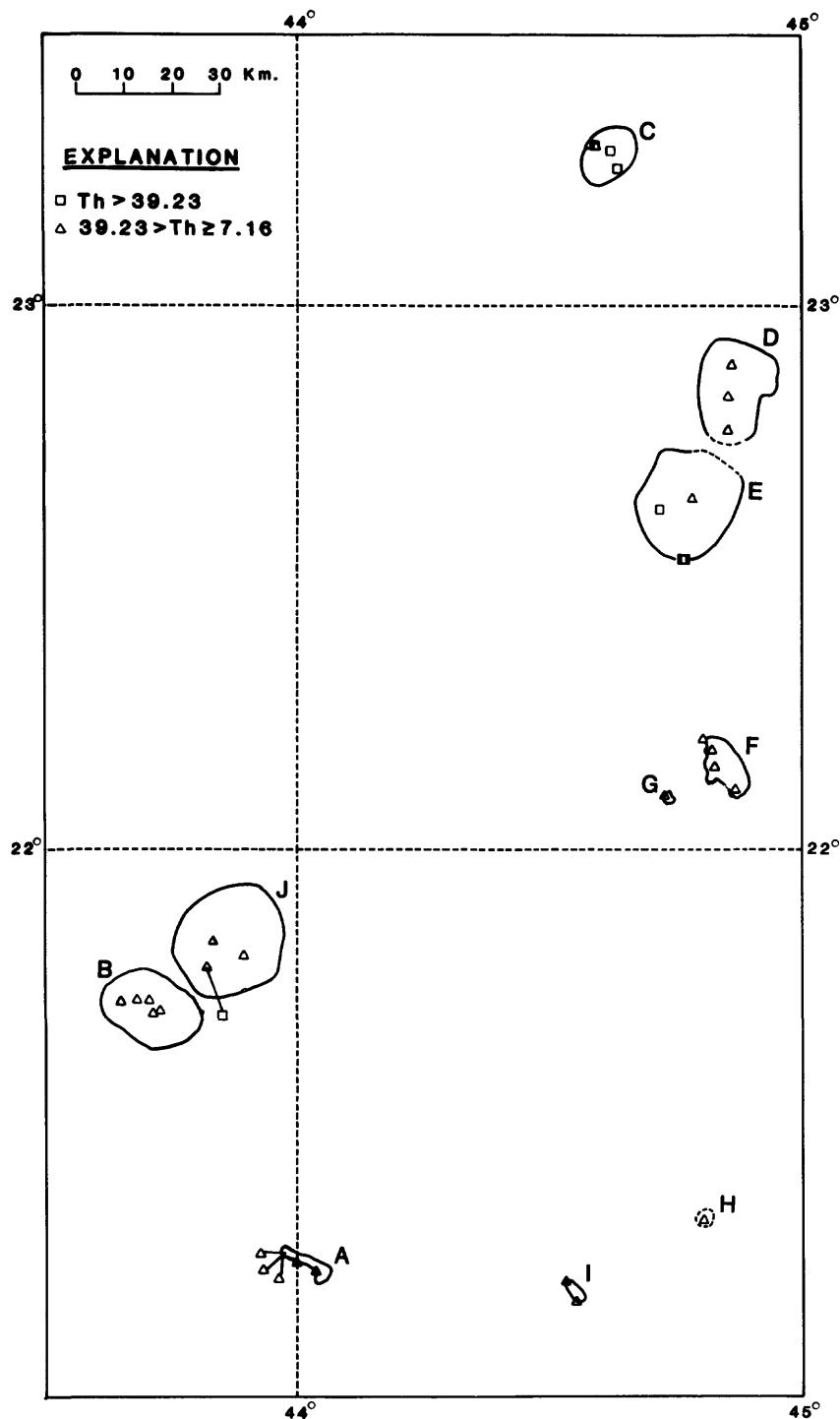


Figure 5.--Map showing areal distribution of thorium values for granite samples from the southeastern part of the Arabian Shield; sample locations and approximate pluton outlines also shown. Symbols represent thorium contents in parts per million relative to the mean and standard deviation of thorium contents in approximately 2,500 granitic samples from the contiguous United States (Stuckless and VanTrump, 1982): triangles, samples having thorium contents within a normal range for granites; squares, anomalously thorium-rich samples; circles, anomalously thorium-poor samples.

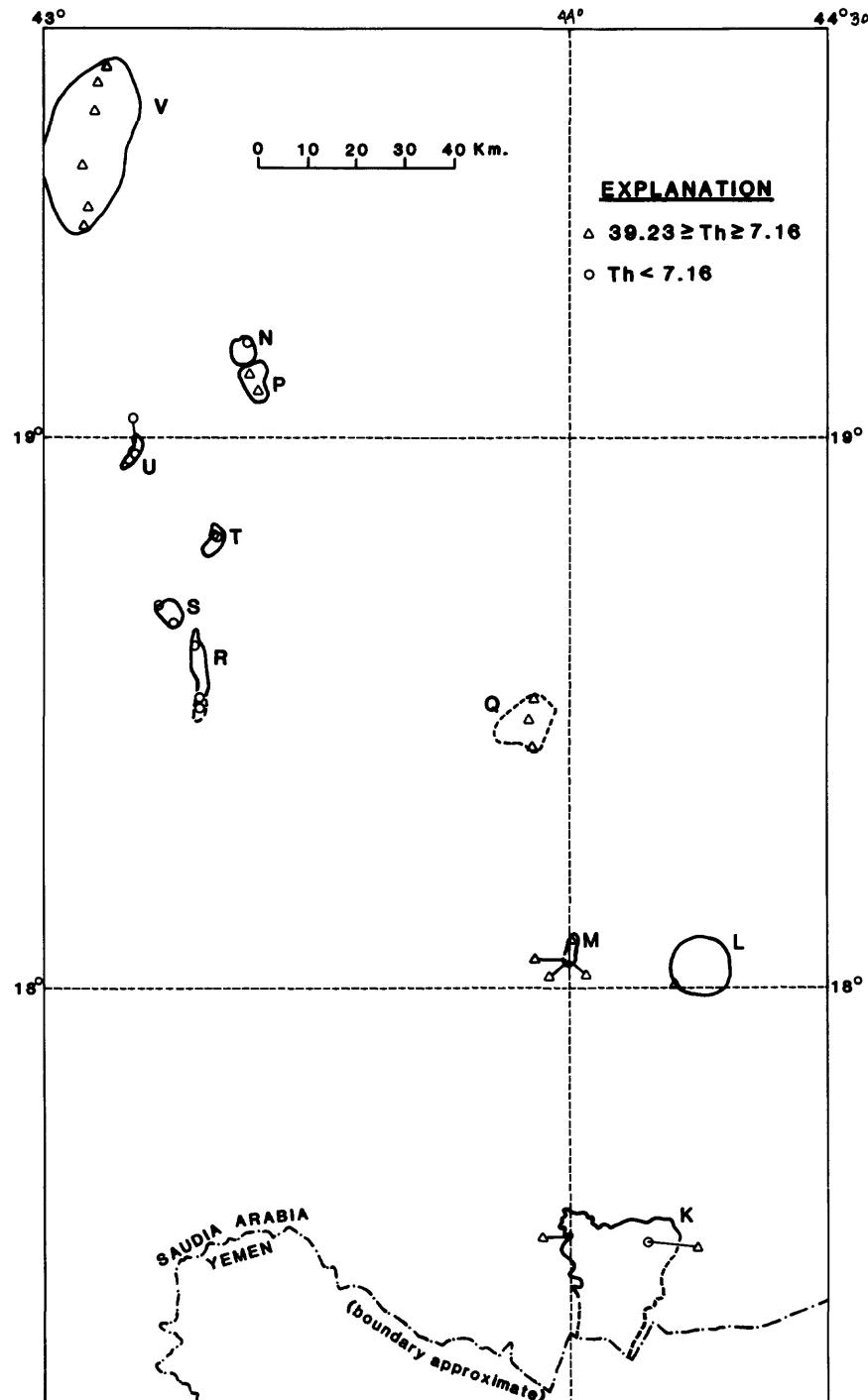


Figure 5.--Continued.

uranium and thorium (0.83, table 5) and the very high correlation coefficient for uranium and radium equivalent uranium (0.98). The latter high correlation coefficient and the general agreement of U and RaeU values within the limits of analytical precision (table 2, fig. 6) indicate that neither uranium nor its relatively short-lived daughter products have been mobilized in the last several thousand years (Stuckless and others, 1977). A map of Th/U ratios (fig. 7) shows that only one of the analyzed samples is anomalously deficient in uranium relative to thorium, and that, on the average, none of the plutons seems to be deficient in uranium relative to thorium.

As noted previously, many of the variations in oxide and trace-element contents can be ascribed to petrogenetic variations such as degree of partial melting or degree of fractional crystallization. It is, therefore, somewhat surprising that a significant percent of the variance in some concentrations can be correlated with sample location. Aluminum, iron, and sodium contents all exhibit a weak, but significant, negative correlation with latitude (table 5). Several trace-element contents are as well or better correlated with latitude (uranium, thorium, and rubidium, table 5). None of the major elements are correlated with longitude to a significant degree; however, several trace elements covary with longitude: uranium, yttrium, thorium, rubidium (table 5), and copper ($r = 0.36$). This is particularly noteworthy in view of the wide range of latitudes ($5^{\circ}34'55''$) relative to the range of longitudes ($1^{\circ}49'03''$). Thus, there is an apparent trend from southwest to northeast, but because this is the general trend of sample locations, the true variations in chemistry could be nearly north-south or east-west.

The observed concentration variations with sample location could be attributed to an increasing degree of magma evolution from southwest to northeast or to a protolithic control in which incompatible-element contents increase from southwest to northeast. The latter interpretation is particularly attractive in view of regional variations in initial-lead isotope compositions noted by Stacey and others (1980) and Stacey and Stoeser (*in press*). These workers attribute the more radiogenic initial leads (in the east) to an older, continental crustal component in the protolith and the less radiogenic initial leads (in the west) to a younger, oceanic crustal source.

The overlap between the areas with chemical data and areas with isotopic data is insufficient to make a direct comparison between the two data sets; however, the northeastern-most area of the current study is close to the area of radiogenic leads. If plutons within the southwestern part of the current study area were shown to contain non-

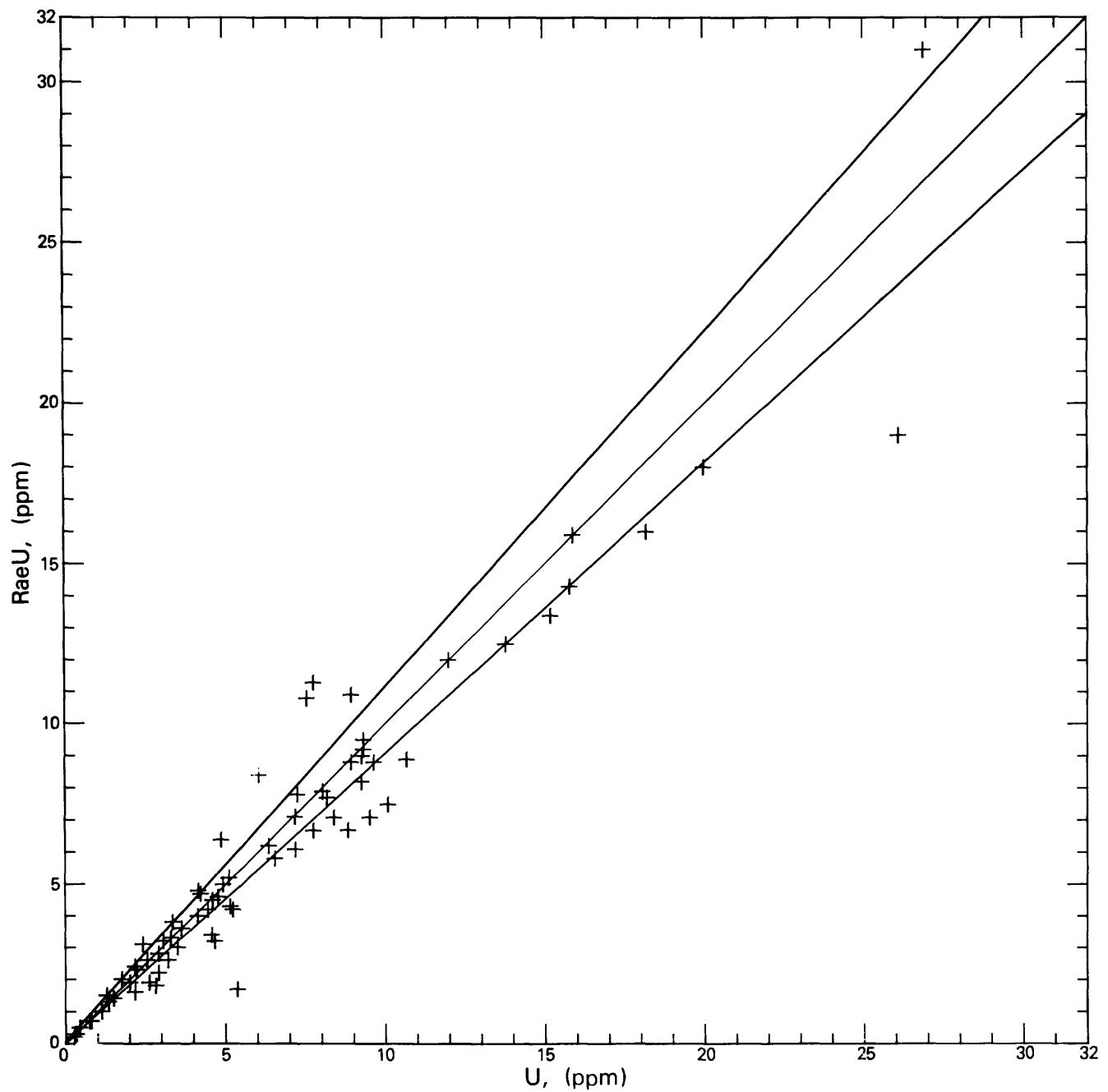


Figure 6.--Diagram showing plot of uranium (U) versus radium equivalent uranium (RaeU) values for granite samples from the southeastern Arabian Shield. Lines show a 1:1 correlation and ± 10 percent.

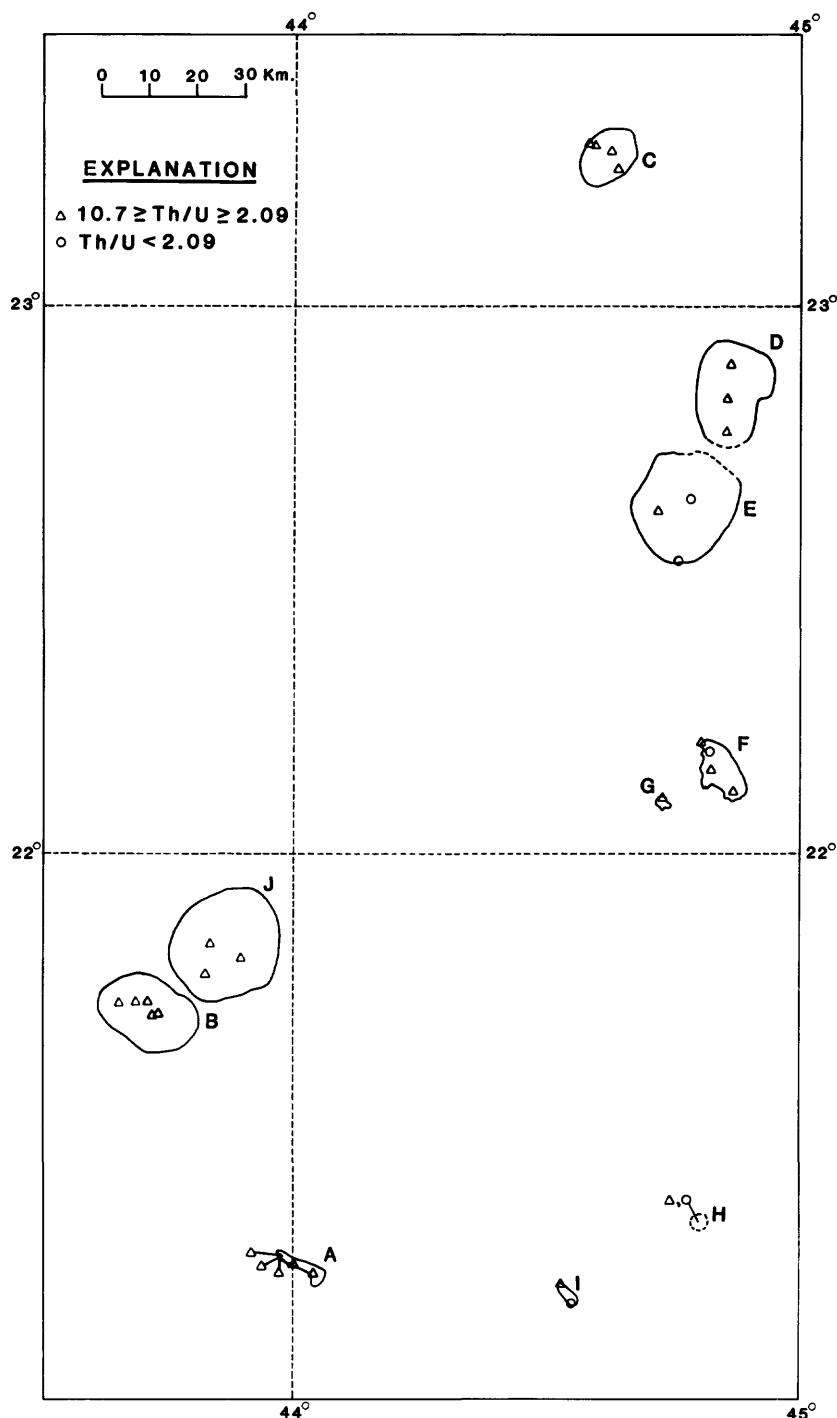


Figure 7.--Map showing areal distribution of Th/U ratios for granite samples from the southeastern part of the Arabian Shield; sample locations and approximate pluton outlines also shown. Symbols represent Th/U ratios relative to the mean and standard deviation of Th/U ratios for approximately 2,500 granitic samples from the contiguous United States (Stuckless and VanTrump, 1982): triangles, samples having Th/U ratios within a normal range for granites; squares, samples having anomalously high Th/U ratios; circles, samples having anomalously low Th/U ratios.

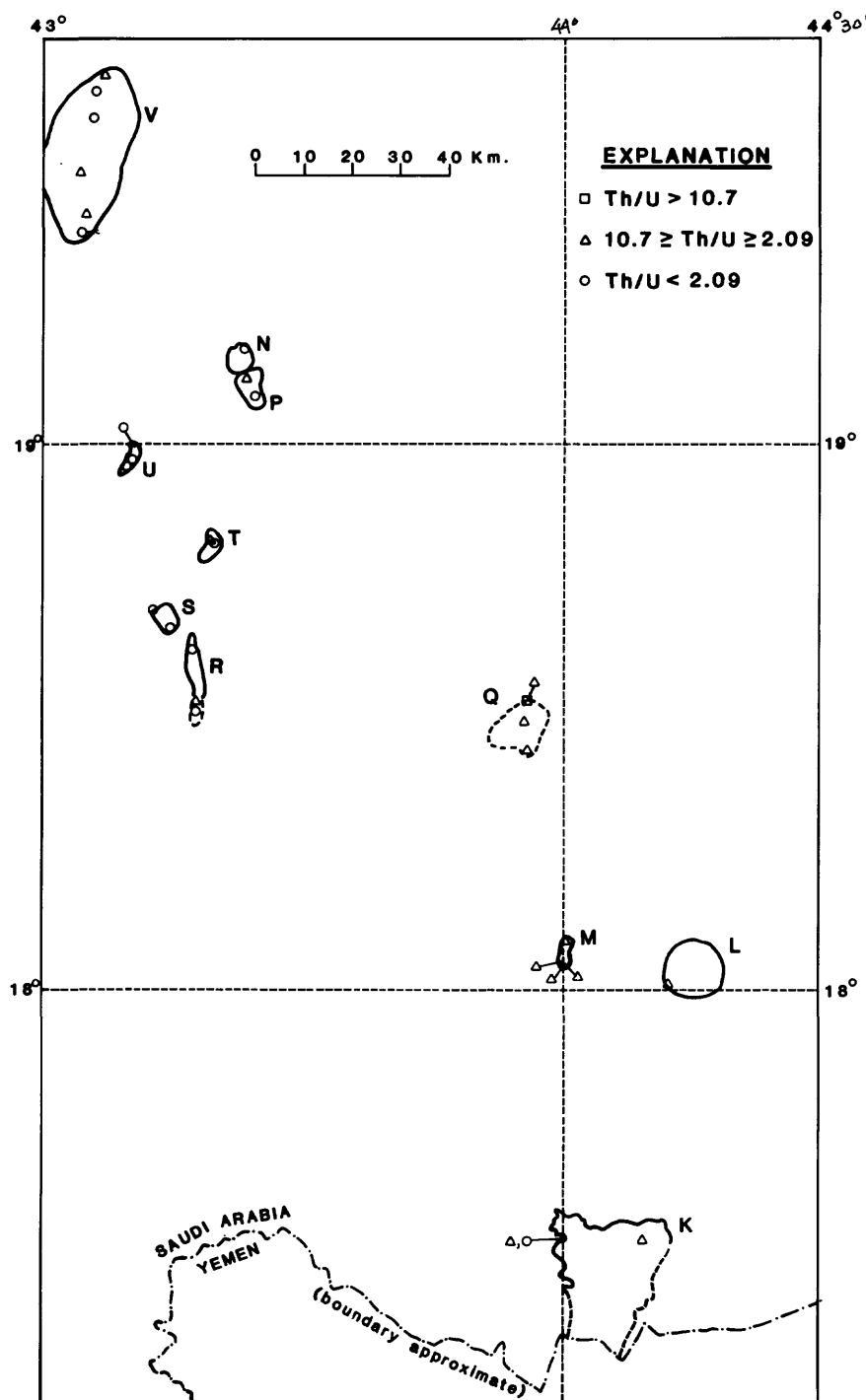


Figure 7.--Continued.

radiogenic leads, the chemical data would strongly support the proposed regional variations in protolith. Higher contents of aluminum, iron, and sodium would be expected from oceanic crustal materials and higher contents of incompatible trace elements would be expected from continental crustal materials.

Suggestions for further work

Concentrations of zinc and copper reported in table 2 are distinctly anomalous, but are not sufficiently precise and accurate to permit detailed evaluation. Reanalysis of at least samples from the more anomalous areas by more accurate techniques (such as induction-coupled plasma spectrometry) is strongly suggested as a first step in confirmation of anomalies. This same technique could be used to simultaneously check tin and tungsten contents, which are known to occur in anomalous amounts elsewhere in postorogenic granites of the eastern Arabian Shield (Cole and others, 1981).

The data in table 5 strongly suggest an areal control for anomalous concentrations of trace elements. A quantitative check of this suggestion should be made by some sort of trend surface analysis. Ideally, compatible data from all postorogenic granites of the Arabian Shield should be used. Most forms of trend surface analysis require a uniform distribution of sample points for a given unit of area, and, therefore, some computer programming and data adjustment would be needed before this approach could be used. Furthermore, the variations in contents of trace elements of interest is dependent in part on degree of magma evolution, and, therefore, some sort of normalizing factor would have to be developed to minimize variations caused by this factor. In spite of the difficulties, the trend-surface-analysis approach should identify the most promising areas for mineral deposits.

The trace-element evolution of granitic rocks, and hence potential for generation of ore deposits, is controlled by the mineralogy of the crystallizing magma. This control is especially important for such elements as uranium, rare-earth elements, niobium, tantalum, tin, and tungsten. For example, tin, and to a lesser extent tungsten, deposits are associated with granites that lack sphene and magnetite (Lehman, 1982), whereas uranium may be associated with sphene-, magnetite-, and allanite-bearing granites (Pagel, 1982). In view of the fact that regional chemical anomalies exist, the determination of trace-mineral suites within the granites and examination of mineral chemistry by microprobe analysis could prove useful in the delineation of probable types of ore deposits.

SUMMARY AND CONCLUSIONS

Copper and zinc occur in anomalous amounts in the postorogenic granites of the southeastern Arabian Shield (table 4) as compared to average granite and are especially anomalous at Jabal Kebad, Jabal Bani Bwana, and Jabal al Hassir (table 2). The southeastern part of the Shield may, therefore, be a favorable area for concentrations of these elements. Data gathered during this study do not suggest a particular mode of occurrence.

The concentrations of yttrium, uranium, and to a lesser extent thorium are on the average at background levels in the postorogenic granites of the southeastern Arabian Shield (table 4), but anomalous concentrations do occur in the northeastern part of the area sampled (figs. 4 and 5, table 4). This area forms the end of a southwest-to-northeast trend of increasing trace-element enrichments. This trend may correspond to recently discovered isotopic patterns and may, therefore, reflect a protolithic control. If a protolithic or other predictable regional control can be identified, optimum areas for detailed work can be defined.

Uranium, thorium, yttrium, and the rare-earth elements have similar geochemical behaviors in an igneous environment, but uranium in particular can be separated readily from these elements in the low-temperature, near-surface environment. High correlation coefficients for uranium, thorium, and yttrium contents, uranium and radium-equivalent uranium contents, and the enrichment of all three elements near a pegmatite indicate that economic deposits of these elements or rare-earth elements would most likely be controlled by magmatic processes. It is suggested that pegmatites associated with anomalous postorogenic granites would have the highest economic potential.

Finally, more detailed field work, quantitative geochemistry, and advanced mathematical techniques are all recommended for further work in the region of the southeastern Arabian Shield.

DATA STORAGE

All results obtained in the course of this study are presented in this report; no data files were created. No existing mineral localities were studied and no specific sites were identified for inclusion in the Mineral Occurrence Documentation System (MODS) of the Saudi Arabian Deputy Ministry for Mineral Resources.

REFERENCES CITED

- Anderson, J. L., and Cullers, R. L., 1978, Geochemistry and evolution of the Wolf River Batholith, a late Precambrian rapakivi massif in North Wisconsin, USA: *Precambrian Research*, v. 7, p. 287-324.
- Anderson, J. L., Cullers, R. L., and Van Schmus, W. R., 1980, Anorogenic metaluminous and peraluminous granite plutonism in the mid-Proterozoic of Wisconsin, USA: *Contributions to Mineralogy and Petrology*, v. 74, p. 311-328.
- Anderson, J. L., *in press*, Proterozoic anorogenic granite plutonism of North America: *Geological Society of America Memoir*.
- Barker, F., Millard, H. T., Jr., Hedge, C. E., and O'Neil, J. R., 1976, Pikes Peak batholith; geochemistry of some minor elements and isotopes, and implications for magma genesis: *in* Epis, R. C., and Weimer, R. J., eds., *Professional Contributions of Colorado School of Mines*, no. 8, p. 44-56.
- Bunker, C. M., and Bush, C. A., 1966, Uranium, thorium, and radium analyses by gamma-ray spectrometry (0.184-0.352 million electron volts), *in* Geological Survey Research 1966: U.S. Geological Professional Paper 550-B, p. B176-B181.
- _____, 1967, A comparison of potassium analyses by gamma-ray spectrometry and other techniques, *in* Geological Survey Research 1967: U.S. Geological Survey Professional Paper 575-B, p. B164-B169.
- Cole, J. C., Smith, C. W., and Fenton, M. D., 1981, Preliminary investigation of the Baid al Jimalah tungsten deposit, Kingdom of Saudi Arabia: U.S. Geological Survey Saudi Arabian Mission Technical Record 20 (Interagency Report 377), 26 p.; also, U.S. Geological Survey Open-File Report 81-1223.
- Cullers, R. L., Koch, R. J., and Bickford, M. E., 1981, Chemical evolution of magmas in the Proterozoic terrane of the St. Francois Mountains, southeastern Missouri, Part 2, Trace element data: *Journal of Geophysical Research*, v. 86, p. 10388-10401.
- Doe, B. R., Stuckless, J. S., and Delevaux, M. H., *in press*, The possible bearing of the granite of the UPH deep drill holes, northern Illinois, on the origin of Mississippi Valley ore deposits: *American Geophysical Union Monograph*.

du Bray, E. A., Elliott, J. E., and Stoeser, D. B., 1983, Geochemical reconnaissance of felsic plutonic rocks in the eastern and southeastern Arabian Shield, Kingdom of Saudi Arabia: U.S. Geological Survey Open-File Report, (IR)SA-491.

Felsche, J., and Herrmann, A. G., 1978, Yttrium and lanthanides 39, 57-71, in K. H. Vedepohl, ed., Handbook of Geochemistry, v. 2, no. 5: Berlin, Springer-Verlag, chapter 39, 57-71.

Fleck, R. J., Greenwood, W. R., Hadley, D. G., Anderson, R. E., and Schmidt, D. L., 1980, Rubidium-strontium geochronology and plate-tectonic evolution of the southern part of the Arabian Shield: U.S. Geological Survey Professional Paper 1131, 38 p.

Harris, N. B. W., and Marriner, G. F., 1980, Geochemistry and petrogenesis of a peralkaline granite complex from the Midian Mountains, Saudi Arabia: Lithos, v. 13, p. 325-337.

Krauskopf, K. B., 1967, Introduction to geochemistry: New York, McGraw-Hill, 721 p.

Lehman, Bernd, 1982, Metallogeny of tin: magmatic differentiation versus geochemical heritage: Economic Geology, v. 77, p. 50-59.

Luth, W. C., Jahns, R. H., and Tuttle, O. F., 1964, The granite system at pressures of 4 to 10 kilobars: Journal of Geophysical Research, v. 69, p. 759-773.

Millard, H. T., Jr., 1976, Determinations of uranium and thorium in U.S. Geological Survey standard rocks by the delayed neutron technique, in Flanagan, F.J., ed., Descriptions and analyses of eight new USGS rock standards: U.S. Geological Survey Professional Paper 840, p. 61-65.

Pagel, Maurice, 1982, The mineralogy and geochemistry of uranium, thorium, and rare-earth elements in two radioactive granites of the Vosges, France: Mineralogical Magazine, v. 46, p. 149-161.

Radaïn, A. A. M., Fyfe, W. S., and Kerrich, R., 1981, Origin of peralkaline granites of Saudi Arabia: Contributions to Mineralogy and Petrology, v. 78, p. 358-366.

Shand, S. J., 1951, Eruptive rocks: New York, John Wiley, 488 p.

Stacey, J. S., Doe, B. R., Roberts, R. J., Delevaux, M. H., and Gramlich, J. W., 1980, A lead isotope study of mineralization in the Saudi Arabian Shield: Contributions to Mineralogy and Petrology, v. 74, p. 175-188.

Stacey, J. S., and Stoesser, D. B., *in press*, Distribution of oceanic and continental leads in the Arabian-Nubian Shield: Contributions to Mineralogy and Petrology.

Stoesser, D. B., and Elliott, J. E., 1980, Post-orogenic peralkaline and calc-alkaline granites and associated mineralization of the Arabian Shield, Kingdom of Saudi Arabia, in Evolution and mineralization of the Arabian-Nubian Shield: King Abdulaziz University, Institute of Applied Geology Bulletin 3, v. 4: Pergamon Press, Oxford-New York, p. 1-23.

Streckeisen, A. L., 1973, Plutonic rocks, classification and nomenclature recommended by the IUGS Subcommission on the Systematics of Igneous Rocks: Geotimes, v. 18, p. 26-30.

Stuckless, J. S., Millard, H. T., Jr., Bunker, C. M., Nkomo, I. T., Rosholt, J. N., Bush, C. A., Huffman, Claude, Jr., and Keil, R. L., 1977, A comparison of some analytical techniques for determining uranium, thorium, and potassium in granitic rocks: U.S. Geological Survey Journal of Research, v. 5, p. 83-91.

Stuckless, J. S., and VanTrump,, George, Jr., 1979, A revised version of Graphic Normative Analysis Program (GNAP) with examples of petrologic problem solving: U.S. Geological Survey Open-File Report 79-1237, 115 p.

Stuckless, J. S., and VanTrump,, George, Jr., 1982, A compilation of radioelement concentrations in granitic rocks of the contiguous United States: Proceedings of the IAEA/OECD Symposium on Uranium Exploration Methods, p. 198-208.

Stuckless, J. S., VanTrump, George, Jr., Bunker, C. M., and Bush, C. A.,*in press*, Preliminary report on the geochemistry and uranium favorability of the postorogenic granites of the northeastern Arabian Shield, Kingdom of Saudi Arabia: Saudi Arabian Deputy Ministry for Mineral Resources Bulletin.

Stuckless, J. S., Knight, R. J., Van Trump, G., Jr., and Budahn, J. R., 1982b, Trace-element geochemistry of post-orogenic granites from the northeastern Arabian Shield, Kingdom of Saudi Arabia: Saudi Arabian Deputy Ministry for Mineral Resources Open-File Report USGS-OF-02-91, 34 p.; also, 1983, U.S. Geological Survey Open-File Report 83-287.

Taggart, J. E., Jr., Lichte, F. E., and Wahlberg, J. S., 1982, Methods of analysis of samples using X-ray fluorescence and induction-coupled plasma spectroscopy, in The 1980 eruptions of Mount St. Helens, Washington: U.S. Geological Survey Professional Paper 1250, p. 683-687.

Thornton, C. P., and Tuttle, O. F., 1960, Chemistry of igneous rocks, I. Differentiation index: American Journal of Science, v. 258, p. 664-684.

Tuttle, O. F., and Bowen, N. L., 1958, Origin of granite in the light of experimental studies in the system NaAlSi₃O₈-KAlSi₃O₈-SiO₂-H₂O: Geological Society of America Memoir 74, 153 p.

VanTrump, George, Jr., and Miesch, A. T., 1977, The U.S. Geological Survey RASS-STATPAC system for management and statistical reduction of geochemical data: Computer and Geosciences, v. 3, p. 475-488.

Watson, E. B., 1979, Zircon saturation in felsic liquids: Experimental results and applications to trace element geochemistry: Contributions to Mineralogy and Petrology, v. 70, p. 407-419.

Whitney, J. A., 1975, The effects of pressure, temperature and X_{H₂O} on phase assemblage in four synthetic rock compositions: Journal of Geology, v. 83, p. 1-31.

Wilson, M. R., and Åkerblom, G., 1980, Uranium enriched granites in Sweden: Sveriges Geologiska Undersökning Rapporter och Meddelanden, v. 19, Uppsala, 30 p.